

## Scientific rationale of Saturn's *in situ* exploration

O. Mousis<sup>a</sup>, L. N. Fletcher<sup>b</sup>, J.-P. Lebreton<sup>c,d</sup>, P. Wurze<sup>e</sup>, T. Cavalié<sup>f</sup>, A. Coustenis<sup>d</sup>, R. Courtin<sup>d</sup>, D. Gautier<sup>d</sup>, R. Helled<sup>g</sup>, P. G. J. Irwin<sup>b</sup>, A. D. Morse<sup>h</sup>, N. Nettelmann<sup>i</sup>, B. Marty<sup>j</sup>, P. Rousselot<sup>a</sup>, O. Venot<sup>k</sup>, D. H. Atkinson<sup>l,n</sup>, J. H. Waite<sup>m</sup>, K. R. Reh<sup>n</sup>, A. Simon-Miller<sup>o</sup>, S. Atreya<sup>p</sup>, N. André<sup>q</sup>, M. Blanc<sup>q</sup>, I. A. Dagleis<sup>r</sup>, G. Fischer<sup>s</sup>, W. D. Geppert<sup>t</sup>, T. Guillot<sup>u</sup>, M. M. Hedman<sup>v</sup>, R. Hueso<sup>w</sup>, E. Lellouch<sup>d</sup>, J. I. Lunine<sup>x</sup>, C. D. Murray<sup>y</sup>, J. O'Donoghue<sup>z</sup>, M. Rengel<sup>f</sup>, A. Sánchez-Lavega<sup>w</sup>, F.-X. Schmider<sup>u</sup>, A. Spiga<sup>aa</sup>, T. Spilker<sup>ab</sup>, J.-M. Petit<sup>a</sup>, M. S. Tiscareno<sup>x</sup>, M. Ali-Dib<sup>a</sup>, K. Altwegg<sup>e</sup>, A. Bouquet<sup>a,m</sup>, C. Briois<sup>c</sup>, T. Fouchet<sup>d</sup>, S. Guerlet<sup>aa</sup>, T. Kostiuk<sup>o</sup>, D. Lebleu<sup>ac</sup>, R. Moreno<sup>d</sup>, G. S. Orton<sup>n</sup>, J. Poncy<sup>ac</sup>

<sup>a</sup> *Université de Franche-Comté, Institut UTINAM, CNRS/INSU, UMR 6213, Observatoire des Sciences de l'Univers de Besançon, France*

<sup>b</sup> *Atmospheric, Oceanic & Planetary Physics, Department of Physics, University of Oxford, Clarendon Laboratory, Parks Road, Oxford OX1 3PU, UK*

<sup>c</sup> *LPC2E, CNRS-Université d'Orléans, 3a Avenue de la Recherche Scientifique, 45071 Orléans Cedex 2, France*

<sup>d</sup> *LESIA, Observatoire de Paris, CNRS, UPMC, Univ. Paris-Diderot, 5, place Jules Janssen, F-92195 Meudon Cedex*

<sup>e</sup> *Space Science & Planetology, Physics Institute, University of Bern, Sidlerstrasse 5, 3012 Bern, Switzerland*

<sup>f</sup> *Max-Planck-Institut für Sonnensystemforschung, Max-Planck-Str. 2, 37191 Katlenburg-Lindau, Germany*

<sup>g</sup> *Department of Geophysics, Atmospheric and Planetary Sciences, Tel-Aviv University, Tel-Aviv, Israel*

<sup>h</sup> *Planetary and Space Sciences, Department of Physics, The Open University, Walton Hall, Milton Keynes MK7 6AA, UK*

<sup>i</sup> *Institute for Physics, University of Rostock, 18051 Rostock, Germany*

<sup>j</sup> *CRPG-CNRS, Nancy-Université, 15 rue Notre Dame des Pauvres, 54501 Vandoeuvre-ls-Nancy, France*

<sup>k</sup> *Instituut voor Sterrenkunde, Katholieke Universiteit Leuven, Celestijnenlaan 200D, 3001 Leuven, Belgium*

<sup>l</sup> *Department of Electrical and Computer Engineering, University of Idaho, Moscow ID 83844-1023, USA*

<sup>m</sup> *Southwest Research Institute (SwRI), 6220 Culebra Road, San Antonio, TX 78228, USA*

<sup>n</sup> *Jet Propulsion Laboratory, California Institute of Technology, 4800 Oak Grove Dr., Pasadena, CA 91109, USA*

---

*Email address: olivier.mousis@obs-besancon.fr (O. Mousis)*

- <sup>o</sup>NASA Goddard Space Flight Center, Code 690, Greenbelt, MD 20771, USA
- <sup>p</sup>Department of Atmospheric, Oceanic, and Space Sciences, University of Michigan, Ann Arbor, MI 48109-2143, USA
- <sup>q</sup> Institut de Recherche en Astrophysique et Planétologie (IRAP), CNRS/Université Toulouse III (UMR 5277), 9, avenue du Colonel Roche, BP 44346, 31028 Toulouse Cedex 4, France, France
- <sup>r</sup>University of Athens, Department of Physics, Panepistimioupoli Zografou, 15784 Athens, Greece
- <sup>s</sup>Space Research Institute, Austrian Academy of Sciences, Schmiedlstrasse 6, A-8042 Graz, Austria
- <sup>t</sup>Stockholm University Astrobiology Centre, Department of Physics, AlbaNova, Stockholm University/Stockholms universitet, Roslagstullbacken 21, S-10691 Stockholm, Sweden/Sverige
- <sup>u</sup>Observatoire de la Côte d'Azur, Laboratoire Lagrange, BP 4229, 06304 Nice cedex 4, France
- <sup>v</sup>Department of Astronomy, Indiana University, Bloomington, Indiana 47405, USA
- <sup>w</sup>Departamento Física Aplicada I, Universidad del País Vasco UPV/EHU, ETS Ingeniería, Alameda Urquijo s/n, 48013 Bilbao, Spain
- Unidad Asociada Grupo Ciencias Planetarias UPV/EHU-IAA(CSIC), 48013 Bilbao, Spain
- <sup>x</sup>Center for Radiophysics and Space Research, Space Sciences Building, Cornell University, Ithaca, NY 14853, USA
- <sup>y</sup>School of Physics and Astronomy, Queen Mary University of London, Mile End Road, London E1 4NS, UK
- <sup>z</sup>Department of Physics and Astronomy, University of Leicester, Leicester LE1 7RH, UK
- <sup>aa</sup>Laboratoire de Météorologie Dynamique, Université Pierre et Marie Curie, Institut Pierre Simon Laplace, Paris, France
- <sup>ab</sup>Solar System Science & Exploration, Monrovia, USA
- <sup>ac</sup>Thales Alenia Space, Cannes, France

---

## Abstract

Remote sensing observations meet some limitations when used to study the bulk atmospheric composition of the giant planets of our solar system. A remarkable example of the superiority of *in situ* probe measurements is illustrated by the exploration of Jupiter, where key measurements such as the determination of the noble gases abundances and the precise measurement of the helium mixing ratio have only been made available through *in situ*

measurements by the Galileo probe. This paper describes the main scientific goals to be addressed by the future *in situ* exploration of Saturn placing the Galileo probe exploration of Jupiter in a broader context and before the future probe exploration of the more remote ice giants. *In situ* exploration of Saturn's atmosphere addresses two broad themes that are discussed throughout this paper: first, the formation history of our solar system and second, the processes at play in planetary atmospheres. In this context, we detail the reasons why measurements of Saturn's bulk elemental and isotopic composition would place important constraints on the volatile reservoirs in the protosolar nebula. We also show that the *in situ* measurement of CO (or any other disequilibrium species that is depleted by reaction with water) in Saturn's upper troposphere would constrain its bulk O/H ratio. We compare predictions of Jupiter and Saturn's bulk compositions from different formation scenarios, and highlight the key measurements required to distinguish competing theories to shed light on giant planet formation as a common process in planetary systems with potential applications to most extrasolar systems. *In situ* measurements of Saturn's stratospheric and tropospheric dynamics, chemistry and cloud-forming processes will provide access to phenomena unreachable to remote sensing studies. Different mission architectures are envisaged, which would benefit from strong international collaborations, all based on an entry probe that would descend through Saturn's stratosphere and troposphere under parachute down to a minimum of 10 bars of atmospheric pressure. We finally discuss the science payload required on a Saturn probe to match the measurement requirements.

*Keywords:* Entry probe, Saturn atmosphere, giant planet formation, solar system formation, *in situ* measurements, elemental and isotopic composition

---

## 1. Introduction

Giant planets contain most of the mass and the angular momentum of our planetary system and must have played a significant role in shaping its large scale architecture and evolution, including that of the smaller, inner worlds (Gomes et al., 2005). Furthermore, the formation of the giant planets affected the timing and efficiency of volatile delivery to the Earth and other terrestrial planets (Chambers and Wetherill, 2001). Therefore, understanding giant planet formation is essential for understanding the origin and evolution of the Earth and other potentially-habitable environments throughout our solar system. The origin of the giant planets, their influence on planetary system architectures, and the plethora of physical and chemical processes at work within their atmospheres, make them crucial destinations for future exploration. Because Jupiter and Saturn have massive envelopes essentially composed of hydrogen and helium and (possibly) a relatively small core, they are called gas giants. Meanwhile, Uranus and Neptune also contain hydrogen and helium atmospheres but, unlike Jupiter and Saturn, their H<sub>2</sub> and He mass fractions are smaller (5 to 20%). They are called ice giants because their density is consistent with the presence of a significant fraction of ices/rocks in their interiors. Despite this apparent grouping into two classes of giant planets, the four giant planets likely exist on a continuum, each a product of the particular characteristics of their formation environment. Comparative planetology of the four giants in the solar system is therefore essential to

reveal the potential formational, migrational, and evolutionary processes at work during the early evolution of the early solar nebula.

Much of our understanding of the origin and evolution of the outer planets comes from remote sensing by necessity. However, the efficiency of this technique has limitations when used to study the bulk atmospheric composition that is crucial to the understanding of planetary origin, namely due to degeneracies between the effects of temperatures, clouds and abundances on the emergent spectra, but also due to the limited vertical resolution. In addition, many of the most common elements are locked away in a condensed phase in the upper troposphere, hiding the main volatile reservoir from the reaches of remote sensing. It is only by penetrating below the “visible” weather layer that we can sample the deeper troposphere where those most common elements are well mixed. A remarkable example of the superiority of *in situ* probe measurements is illustrated by the exploration of Jupiter, where key measurements such as the determination of the noble gases abundances and the precise measurement of the helium mixing ratio have only been possible through *in situ* measurements by the Galileo probe (Owen et al., 1999).

The Galileo probe measurements provided new insights into the formation of the solar system. For instance, they revealed the unexpected enrichments of Ar, Kr and Xe with respect to their solar abundances, which suggested that the planet accreted icy planetesimals formed at temperatures possibly as low as 20–30 K to allow the trapping of these noble gases. Another remarkable result was the determination of the Jovian helium abundance using a dedicated instrument aboard the Galileo probe (von Zahn et al., 1998) with an accuracy of 2%. Such an accuracy on the He/H<sub>2</sub> ratio is impossible to derive

from remote sensing, irrespective of the giant planet being considered, and yet precise knowledge of this ratio is crucial for the modelling of giant planet interiors and thermal evolution. The Voyager mission has already shown that these ratios are far from being identical, which presumably results from slight differences in their histories at different heliocentric distances. An important result also obtained by the mass spectrometer onboard the Galileo probe was the determination of the  $^{14}\text{N}/^{15}\text{N}$  ratio, which suggested that nitrogen present in Jupiter today originated from the solar nebula essentially in the form of  $\text{N}_2$  (Owen et al., 2001). The Galileo science payload unfortunately could not probe to pressure levels deeper than 22 bars, precluding the determination of the  $\text{H}_2\text{O}$  abundance at levels representative of the bulk oxygen enrichment of the planet. Furthermore, the probe descended into a region depleted in volatiles and gases by unusual “hot spot” meteorology (Orton et al., 1998; Wong et al., 2004), and therefore its measurements are unlikely to represent the bulk planetary composition. Nevertheless, the Galileo probe measurements were a giant step forward in our understanding of Jupiter. However, with only a single example of a giant planet measurement, one must wonder whether from the measured pattern of elemental and isotopic enrichments, the chemical inventory and formation processes at work in our solar system are truly understood. *In situ* exploration of giant planets is the only way to firmly characterize the planet compositions in the solar system. In this context, a Saturn probe is the next natural step beyond Galileo’s *in situ* exploration of Jupiter, the remote investigation of its interior and gravity field by the JUNO mission, and the Cassini spacecraft’s orbital reconnaissance of Saturn.

*In situ* exploration of Saturn’s atmosphere addresses two broad themes. First, the formation history of our solar system and second, the processes at play in planetary atmospheres. Both of these themes are discussed throughout this paper. Both themes have relevance far beyond the leap in understanding gained about an individual giant planet: the stochastic and positional variances produced within the solar nebula, the depth of the zonal winds, the propagation of atmospheric waves, the formation of clouds and hazes and disequilibrium processes of photochemistry and vertical mixing are common to all planetary atmospheres, from terrestrial planets to gas and ice giants and from brown dwarfs to hot exoplanets.

This paper describes the main scientific goals to be addressed by the future *in situ* exploration of Saturn placing the Galileo probe exploration of Jupiter in a broader context and before the future *in situ* exploration of the more remote ice giants. These goals will become the primary objectives listed in the forthcoming Saturn probe proposals that we intent to submit in response to future opportunities within both ESA and NASA. Section 2 is devoted to a comparison between known elemental and isotopic compositions of Saturn and Jupiter. We describe the different formation scenarios that have been proposed to explain Jupiter’s composition and discuss the key measurements at Saturn that would allow disentangling these interpretations. We also demonstrate that the *in situ* measurement of CO (or any other disequilibrium species that is depleted by reaction with water) at Saturn could place limits on its bulk O/H ratio. In Section 3, we discuss the motivation for the *in situ* observation of the atmospheric processes (dynamics, chemistry and cloud formation) at work in Saturn’s atmosphere. Section 4 is dedicated to a

short description of the mission designs that can be envisaged. In Section 5, we provide a description of high-level specifications for the science payload. Conclusions are given in Section 6.

## **2. Elemental and Isotopic Composition as a Window on Saturn's Formation**

The giant planets in the solar system formed 4.55 Gyr ago from the same material that engendered the Sun and the entire solar system. The envelopes of giant planets are dominated by hydrogen and helium, the two most abundant elements in the Universe. Protoplanetary disks, composed of gas and dust, are almost ubiquitous when stars form, but their typical lifetimes do not exceed a few million years. This implies that the gas giants Jupiter and Saturn had to form rapidly to capture their hydrogen and helium envelopes, more rapidly than the tens of millions of years needed for terrestrial planets to reach their present masses. Due to formation at fairly large radial distances from the Sun, where the solid surface density is low, the ice giants Uranus and Neptune had longer formation timescales (slow growth rates) and did not manage to capture large amounts of hydrogen and helium before the disk gas dissipated. As a result, the masses of their gaseous envelopes are small compared to their ice/rock cores.

A comparative study of the properties of these giant planets thus gives information on spatial gradients in the physical/chemical properties of the solar nebula as well as on stochastic effects that led to the formation of the solar system. Data on the composition and structure of the giant planets, which hold more than 95% of the non-solar mass of the solar system, remain



scarce, despite the importance of such knowledge. The formation of giant planets is now largely thought to have taken place via the core accretion model in which a dense core is first formed by accretion and the hydrogen-helium envelope is captured after a critical mass is reached (Mizuno, 1980; Pollack et al., 1996). When the possibility of planet migration is included (Lin and Papaloizou, 1986; Ward, 1997), such a model can explain the orbital properties of exoplanets, although lots of unresolved issues remain (Ida and Lin, 2004; Mordasini et al., 2012). However, an alternative scenario for the formation of giant planets is the disk instability model (Boss, 1997, 2001), in which the giant planets form from the direct contraction of a gas clump resulting from local gravitational instability in the disk. In principle, measurements of atmospheric bulk elemental enrichments and isotopic ratios would help us to distinguish between these competing formation scenarios.

Formation and evolution models indicate that the total mass of heavy elements present in Jupiter may be as high as  $42 M_{\oplus}$ , whereas the mass of the core is estimated to range between 0 and  $13 M_{\oplus}$  (Saumon and Guillot, 2004). In the case of Saturn, the mass of heavy elements can be as large as  $35 M_{\oplus}$  with a mass varying between 0 and  $10 M_{\oplus}$  in the envelope and the core mass ranging between 0 and  $20 M_{\oplus}$  (Helled and Guillot, 2013). Direct access to heavy materials within giant planet cores to constrain these models is impossible, so we must use the composition of the well-mixed troposphere to infer the properties of the deep interiors. It is difficult for remote sounding to provide the necessary information because of a lack of sensitivity to the atmospheric compositions beneath the cloudy, turbulent and chaotic weather layer. These questions must be addressed by *in situ* exploration, even if the

NASA JUNO mission will try to address them remotely.

The availability of planetary building blocks (metals, oxides, silicates, ices) is expected to vary with position within the original nebula, from refractories in the warm inner nebula to a variety of ices of water,  $\text{CH}_4$ ,  $\text{CO}$ ,  $\text{NH}_3$ ,  $\text{N}_2$  and other simple molecules in the cold outer nebula. Turbulent radial mixing, and the evolution of the pressure-temperature gradient in the disk could have led to distinct regions where some species dominated over others (e.g., the water ice snowline or  $\text{N}_2$  over  $\text{NH}_3$ ). Furthermore, both inward and outward migration of the giants during their evolution could have provided access to different material reservoirs at different epochs. A giant planet's bulk composition therefore depends on the timing and location of planet formation, subsequent migration and the delivery mechanisms for the heavier elements. By measuring a giant planet's chemical inventory, and contrasting it with measurements of (i) other giant planets, (ii) primitive materials found in comets and asteroids, and (iii) the elemental abundances of our parent star and the local interstellar medium, we can reveal much about the conditions at work during the formation of our planetary system.

It should be noted, however, that when atmospheric measurements are used to infer the planetary composition and reveal information on the planet's origin, one has to *assume* that the atmospheric composition represents the bulk planetary composition. This is a fairly good assumption if the measurements probe the convective region, and if the planet is fully convective. Within a fully convective planet the materials are expected to be homogeneously mixed, and therefore, we do not expect large differences in composition with depth. However, if the planet is not fully convective and homoge-

neously mixed, the information of its atmospheric composition cannot solely be used to infer the bulk composition.

In the case of Saturn (as well as Jupiter) compositional inhomogeneities can be the outcome of the formation process (e.g. [Pollack et al., 1996](#)) and/or the erosion of a primordial core that could mix with the surrounding metallic hydrogen ([Guillot, 2004](#); [Wilson and Militzer, 2011, 2012](#)). In addition, it is possible that double diffusive convection occurs in the interiors of giant planets (e.g. [Leconte and Chabrier, 2012, 2013](#)). If a molecular weight gradient is maintained throughout the planetary envelope, double-diffusive convection would take place, and the thermal structure would be very different from the one that is generally assumed using adiabatic (i.e., fully convective) models, with much higher center temperatures and a larger fraction of heavy elements. In this case, the planetary composition can vary substantially with depth and therefore, a measured composition of the envelope would not represent the overall composition. While standard interior models of Saturn assumed three layers and similar constraints in terms of the helium to hydrogen ratio, they can differ in the assumption on the distribution of heavy elements within the planetary envelope. While Guillot and collaborators (e.g. [Saumon and Guillot, 2004](#); [Helled and Guillot, 2013](#)) assume homogeneous distribution of heavy elements apart from helium, which is depleted in the outer envelope due to helium rain (a process that seems to be related also to neon depletion), interior structure models by Nettelmann and collaborators ([Fortney and Nettelmann, 2010](#); [Nettelmann et al., 2013](#)) allow the abundance of heavy elements to be discontinuous between the molecular and the metallic envelope. At present, it is not clear whether there should be a dis-

continuity in the composition of heavy elements, and this question remains open.

### 2.1. *Jupiter and Saturn's Composition*

The abundances and isotopic ratios of most significant volatiles measured at Jupiter and Saturn are given in Tables 1 and 2. We refer the reader to the papers of [Atreya et al. \(2003\)](#), [Teanby et al. \(2006\)](#) and [Fletcher et al. \(2012\)](#) for a more exhaustive list of disequilibrium species identified (or for other minor species presumably identified) in Jupiter's and Saturn's atmospheres. Only upper limits on the abundances of hydrogen halides have been derived from the remote detection of these species in Saturn's atmosphere, implying the need of a probe to get improved *in situ* measurements.

The abundances of CH<sub>4</sub>, NH<sub>3</sub>, H<sub>2</sub>O, H<sub>2</sub>S, Ne, Ar, Kr and Xe have been measured by the Galileo Probe Mass Spectrometer (GPMS) in Jupiter's atmosphere ([Mahaffy et al., 2000](#); [Wong et al., 2004](#)). The value of H<sub>2</sub>O abundance reported for Jupiter in Table 1 corresponds to the deepest measurement made by the probe (at 17.6–20.9 bar) and is probably much smaller than the planet's bulk water abundance, which remains unknown ([Atreya et al., 2003](#); [Wong et al., 2004](#)). The Juno mission, which will arrive at Jupiter in 2016, may provide an estimate of the tropospheric O/H ratio. The He abundance in Jupiter has also been measured *in situ* by a Jamin-Mascart interferometer aboard the Galileo probe (Helium Abundance Detector; hereafter HAD) with a better accuracy level than the GPMS instrument ([von Zahn et al., 1998](#)). PH<sub>3</sub> is the only species of our list of Jupiter measurements whose abundance has been determined remotely by the Cassini Composite Infrared Spectrometer (CIRS) during the spacecraft 2000–2001 encounter ([Fletcher et](#)

al., 2009a).  $\text{PH}_3$  is a disequilibrium species at its sampling level in Jupiter's atmosphere (see Sec. 3), implying that its bulk abundance is probably higher than the measured one. Isotopic measurements presented for Jupiter in Table 2 have also been performed by the GPMS instrument aboard the Galileo probe (Niemann et al., 1996, 1998; Mahaffy et al., 2000; Atreya et al., 2003; Wong et al., 2004).

In the case of Saturn, only the abundances of  $\text{CH}_4$ ,  $\text{PH}_3$ ,  $\text{NH}_3$  and  $\text{H}_2\text{O}$ , and indirectly that of  $\text{H}_2\text{S}$ , have been measured. The abundance of  $\text{CH}_4$  has been determined from the analysis of high spectral resolution observations from CIRS (Fletcher et al., 2009b). Similarly to Jupiter,  $\text{PH}_3$  has been determined remotely in Saturn from Cassini/CIRS observations at  $10 \mu\text{m}$  (Fletcher et al., 2009a). Other measurements of  $\text{PH}_3$  have been made from ground based observations at  $5 \mu\text{m}$  (de Graauw et al., 1997), but the spectral line data at these wavelengths is less robust and accurate than those at  $10 \mu\text{m}$ . There is also a degeneracy with the location, extent, opacity of Saturn's clouds at  $5 \mu\text{m}$  which is not apparent at  $10 \mu\text{m}$ . Moreover, considering the fact that there is also terrestrial contamination in the  $5 \mu\text{m}$  window for groundbased observations and that the scattered sunlight may contribute at  $5 \mu\text{m}$ , this leads us to believe that the data at  $10 \mu\text{m}$  are more reliable. Interestingly, we note that  $\text{PH}_3$  is easier to detect on Saturn compared to Jupiter because this molecule dominates the upper tropospheric chemistry and ammonia is locked away at deeper levels. The  $\text{NH}_3$  abundance is taken from the range of values derived by Fletcher et al. (2011) who analyzed Saturn's tropospheric composition from Cassini/VIMS  $4.6\text{--}5.1 \mu\text{m}$  thermal emission spectroscopy. This determination is probably more reliable than those made

in the microwave domain because of the absence of spectral lines at these wavelengths (Briggs and Sackett, 1989; Laraia et al., 2013). Tropospheric H<sub>2</sub>O has been inferred in Saturn via the Short Wavelength Spectrometer Instrument onboard the Infrared Space Observatory (ISO-SWS) (de Graauw et al., 1997). However, H<sub>2</sub>O is unsaturated at this altitude ( $\sim 3$  bar level), implying that its bulk abundance is higher than the measured one. The H<sub>2</sub>S abundance is quoted from the indirect determination of Briggs and Sackett (1989) who investigated the influence of models of NH<sub>3</sub>-H<sub>2</sub>S-H<sub>2</sub>O cloud decks on Saturn’s atmospheric opacity at microwave wavelengths. The He abundance in Saturn’s atmosphere derives from a reanalysis of Voyager’s infrared spectrometer (IRIS) measurements (Conrath and Gautier, 2000). The only isotopic ratios measured in Saturn are D/H in H<sub>2</sub> (determination from ISO-SWS, Lellouch et al., 2001) and <sup>12</sup>C/<sup>13</sup>C in CH<sub>4</sub> (Cassini/CIRS observations, Fletcher et al., 2009b).

Table 3 summarizes the enrichments in volatiles relative to protosolar values observed in Jupiter and Saturn. Note that protosolar abundances are different from present-day solar photospheric abundances because elements heavier than He are settling out of the photosphere over time. This mechanism leads to a fractionation of heavy elements relative to hydrogen in the solar photosphere, requiring the use of correction terms to retrieve the protosolar abundances (Asplund et al., 2009). For the sake of information, the protosolar elemental abundances used in our calculations are detailed in Table 4. C, N, P, S, Ar, Kr and Xe are all found enriched by a factor  $\sim 2$  to 4 in Jupiter. On the other hand, C, N and P (the only heavy elements *a priori* reliably measured) are found enriched by factors of  $\sim 10$ , 0.5–5 and 11.5 in

Saturn. Helium is depleted compared to protosolar values in the two giants because of its condensation into droplets that “rain out” in the giant planets deep interiors (Stevenson and Salpeter, 1977a,b; Fortney and Hubbard, 2003). The solution of neon in those droplets (Wilson and Militzer, 2010) would also explain its apparent depletion in Jupiter but a similar measurement has never been possible on Saturn. As mentioned above, oxygen is also depleted compared to protosolar in the Jovian atmosphere but this measurement results from the fact that the Galileo probe entry site was an unusually dry meteorological system. As a result, the probe did not measure the deep, well-mixed water mixing ratio (Wong et al., 2004), which is predicted to be supersolar (Stevenson and Lunine, 1988; Gautier et al., 2001; Hersant et al., 2004; Alibert et al., 2005; Mousis et al., 2009, 2012).

## 2.2. Indirect Determination of Saturn’s O/H Ratio

One of the main objectives of Saturn’s *in situ* exploration is the measurement of the H<sub>2</sub>O abundance. However, depending on the O/H elemental enrichment (Atreya et al., 1999), H<sub>2</sub>O is predicted to condense in the 12.6–21 bar range and may remain out of reach for the probe we consider in this paper that would be limited to ~10 bar (see Sec. 4). Several disequilibrium species, like CO, can provide useful constraints on Saturn’s deep H<sub>2</sub>O abundance. The upper tropospheric mole fraction of CO is representative of the H<sub>2</sub>O abundance in the deep hot troposphere, where the two species are in thermochemical equilibrium (Fegley and Lodders, 1994). It is thus possible to derive the deep H<sub>2</sub>O abundance from CO observations using the “quench level” approximation (e.g., Bézard et al. 2002), or more rigorously using comprehensive thermochemical models (e.g., Visscher et al. 2010 and

Cavalié et al. 2014).

We have adapted the model of Venot et al. (2012) to Saturn’s troposphere to assess the relevance of measuring CO with an *in situ* probe. The thermochemical kinetic network comes from the engine industry and was thoroughly validated for high temperatures and pressures. The tropospheric thermal profile has been constructed from a recent retrieval of the latitudinally-resolved  $T(P)$  structure representing a mean of Cassini’s prime mission (Fletcher et al., 2009b). We used the nominal mixing ratios from Table 1 for He and CH<sub>4</sub>, and adopted an upper limit of  $10^{-9}$  for CO (Cavalié et al., 2009). We have assumed a vertically constant eddy mixing coefficient  $K_{zz}$  ranging from  $10^8$  to  $10^9$  cm<sup>2</sup>·s<sup>-1</sup> (Visscher et al., 2010). With  $K_{zz}=10^8$  cm<sup>2</sup>·s<sup>-1</sup>, the deep atmospheric O/H ratio needs to be 62 times the protosolar value to reproduce the CO upper limit. With  $K_{zz}=10^9$  cm<sup>2</sup>·s<sup>-1</sup>, the O/H still needs to be 18 times protosolar (see Fig. 1), i.e., still much higher than Saturn’s C/H ratio (9.9 times protosolar) but remains within the range of values predicted from the theory arguing that volatiles formed clathrates and pure condensates in the nebula (see Sec. 2.3.2). If we reversely set O/H ratio to the C/H one, then the most favorable case for a detection of CO ( $K_{zz}=10^9$  cm<sup>2</sup>·s<sup>-1</sup>) gives an upper tropospheric mole fraction of CO of  $4.1 \times 10^{-10}$ . Reaching such a low value will remain very challenging for any ground-based facility. Besides, a complication comes from the fact that the observable CO vertical profile is largely dominated by an external source in the stratosphere (Cavalié et al., 2010).

These results argue in favor of an *in situ* measurement of tropospheric CO with a neutral mass spectrometer as a valuable complement to any attempt to



directly measure the H<sub>2</sub>O abundance. However, CO has a molecular weight very close to that of N<sub>2</sub>. This degeneracy is a serious issue because the N<sub>2</sub> upper tropospheric mole fraction is expected to be around four orders of magnitude higher than the one of CO. A mass spectrometer will therefore need a mass resolution of  $m/\Delta m = 2,500$  to separate CO from N<sub>2</sub> at equal abundance, and about  $m/\Delta m = 15,000$  for the CO and N<sub>2</sub> abundances expected in Saturn’s atmosphere. More generally, any other disequilibrium species that reacts with H<sub>2</sub>O, like PH<sub>3</sub> and SiH<sub>4</sub>, is likely to provide additional constraints on the deep H<sub>2</sub>O abundance of Saturn (Visscher and Fegley, 2005) and it would be desirable to include the combustion reaction schemes of such species (e.g., Twarowski 1995 and Miller 2004) in thermochemical models.

### *2.3. Isotopic Measurements at Saturn*

As shown in Table 2, very little is known today concerning the isotopic ratios in Saturn’s atmosphere. Only D/H (for H<sub>2</sub> and methane) and <sup>12</sup>C/<sup>13</sup>C (for methane) ratios have been measured so far (Lellouch et al., 2001; Fletcher et al., 2009b).

The case of D/H is interesting and would deserve further measurements with smaller errors. Because deuterium is destroyed in stellar interiors and transformed into <sup>3</sup>He, the D/H value presently measured in Jupiter’s atmosphere is estimated to be larger by some 5–10% than the protosolar value. This slight enrichment would have resulted from a mixing of nebular gas with deuterium-rich ices during the planet’s formation, as suggested by Guillot (1999). For Saturn, the contribution of deuterium-rich ices in the present D/H ratio could be higher (25–40%). An accurate measurement of the D/H ratio in Saturn’s atmosphere could provide, consequently, some constraints

on the relative contribution of deuterium-rich ices during the formation of Saturn. Such a constraint is also based on the *a priori* knowledge of the protosolar D/H ratio, which remains relatively uncertain. This ratio is estimated from measurements of  $^3\text{He}/^4\text{He}$  in the solar wind, which is corrected for changes that occurred in the solar corona and chromosphere subsequently to the evolution of the Sun's interior, and to which the primordial  $^3\text{He}/^4\text{He}$  is subtracted. This latter value is estimated from the ratio observed in meteorites or in Jupiter's atmosphere. The measurement of  $^3\text{He}/^4\text{He}$  in Saturn's atmosphere would also complement, consequently, the scientific impact of D/H measurement. In any case the smaller value of D/H measured by [Lelouch et al. \(2001\)](#) in Saturn's atmosphere from infrared spectra obtained by the Infrared Space Observatory (ISO) satellite and the Short Wavelength Spectrometer (SWS) compared to Jupiter's atmosphere ([Niemann et al., 1998](#)) is surprising in the sense that it would suggest a lower relative contribution of deuterium-rich ices in the formation of Saturn compared to Jupiter. These values have, nevertheless, large errors and so far no clear conclusion can be drawn.

The  $^{14}\text{N}/^{15}\text{N}$  ratio presents large variations in the different planetary bodies in which it has been measured and, consequently, remains difficult to interpret. The analysis of Genesis solar wind samples ([Marty et al., 2011](#)) suggests a  $^{14}\text{N}/^{15}\text{N}$  ratio of  $441 \pm 5$ , which agrees with the *in situ* measurements made in Jupiter's atmospheric ammonia ([Fouchet et al., 2000, 2004](#)) which probably comes from primordial  $\text{N}_2$  ([Owen et al., 2001](#)). Terrestrial atmospheric  $\text{N}_2$ , with a value of 272, appears enriched in  $^{15}\text{N}$  compared to Jupiter and similar to the bulk of ratios derived from the analysis of comet

81P/ Wild 2 grains (McKeegan et al., 2006). Measurements performed in Titan's atmosphere, which is dominated by  $N_2$  molecules, lead to  $167.7 \pm 0.6$  and  $147.5 \pm 7.5$  from the Cassini/INMS and Huygens/GCMS data, respectively (Niemann et al., 2010; Mandt et al., 2009). Because of the low abundance of primordial Ar observed by Huygens, it is generally assumed that  $N_2$  is of secondary origin in Titan's atmosphere and that N was delivered in a less volatile form, probably  $NH_3$ . Different mechanisms have been proposed for the conversion of  $NH_3$  to  $N_2$ . Isotopic fractionation may have occurred for nitrogen in Titan's atmosphere but the atmospheric model published by Mandt et al. (2009) suggests that the current  $^{14}N/^{15}N$  ratio observed in  $N_2$  is close to the value acquired by the primordial ammonia of Titan. This statement is supported by the recent measurement of the  $^{14}N/^{15}N$  isotopic ratio in cometary ammonia (Rousselot et al., 2014). This ratio, comprised between 80 and 190, is consistent with the one measured in Titan.

All these measurements suggest that  $N_2$  and  $NH_3$  result from the separation of nitrogen into at least two distinct reservoirs, with a distinct  $^{15}N$  enrichment, which never equilibrated. The reservoir containing  $N_2$  would have a large  $^{14}N/^{15}N$  ratio (like in Jupiter's atmosphere, where the present ammonia is supposed to come from primordial  $N_2$ ) and the one containing  $NH_3$  a much lower value (like in Titan's atmosphere, where the present  $N_2$  could come from primordial ammonia, and in cometary ammonia). In this context measuring  $^{14}N/^{15}N$  in Saturn's atmosphere would be very helpful to get more information about the origin of ammonia in this planet.

The cases of carbon, oxygen and noble gas (Ne, Ar, Kr, and Xe) isotopic ratios are different because they should be representative of their primor-

dial values. Only little variations are observed for the  $^{12}\text{C}/^{13}\text{C}$  ratio in the solar system irrespective of the body and molecule in which it has been measured. This ratio appears compatible with the terrestrial value of 89 (except if isotopic fractionation processes occur, like for methane in Titan, but the influence of these processes on this ratio is small). Table 2 provides the value of 91.8 measured by [Fletcher et al. \(2009b\)](#) in Saturn with the Cassini/CIRS but with large error bars. A new *in situ* measurement of this ratio should be useful to confirm that carbon in Saturn is also representative of the protosolar value (and different from the one present in the local Interstellar Medium (ISM) because  $^{13}\text{C}$  is created in stars). The oxygen isotopic ratios also constitute interesting measurements to be made in Saturn’s atmosphere. The terrestrial  $^{16}\text{O}/^{18}\text{O}$  and  $^{16}\text{O}/^{17}\text{O}$  isotopic ratios are 499 and 2632, respectively ([Asplund et al., 2009](#)). At the high accuracy levels possible with meteorites analysis these ratios present some small variations<sup>1</sup>. Measurements performed for solar system objects like comets, far less accurate, match the terrestrial  $^{16}\text{O}/^{18}\text{O}$  value (with error bars being typically a few tens). However no  $^{16}\text{O}/^{18}\text{O}$  ratio has been yet published for Saturn’s atmosphere. The only  $^{16}\text{O}/^{18}\text{O}$  measurement made so far for a giant planet ([Noll et al., 1995](#)) was obtained from groundbased IR observations in Jupiter’s atmosphere and had a very large uncertainty (1–3 times the terrestrial value).

---

<sup>1</sup>Expressed in  $\delta$  units, which are deviations in part per thousand, they are typically a few units.

#### *2.4. Interpretations of the Volatile Enrichments in Jupiter and Saturn*

Several theories connecting the thermodynamic evolution of the protosolar nebula to the formation conditions of the giant planets have been developed to interpret the volatile enrichments measured in Jupiter and Saturn. The main scenarios proposed in the literature and their predictions for Saturn's composition are summarized below.

##### *2.4.1. Amorphous Ice Scenario*

The model proposed by [Owen et al. \(1999\)](#) is the first attempt to explain of the volatile enrichments measured in Jupiter's atmosphere. In this scenario, volatiles present in Jupiter's atmosphere were first acquired in amorphous ice at temperatures as low as 30 K in the protosolar nebula. This hypothesis is based on the fact that formation scenarios of the protosolar nebula invoke two reservoirs of ices, namely an inner and an outer reservoir, that took part in the production of icy planetesimals. The first reservoir, located within  $\sim 30$  Astronomical Units (AU) of the Sun, contains ices (mostly water ice) originating from the ISM which, due to their proximity to the Sun, were initially vaporized ([Chick and Cassen, 1997](#)). With time, the decrease of temperature and pressure conditions allowed the water in this reservoir to condense at  $\sim 150$  K at nebular pressure conditions in the form of (microscopic) crystalline ice ([Kouchi et al., 1994](#)). The other reservoir, located at larger heliocentric distances, is composed of ices originating from the ISM that did not vaporize when entering into the disk. In this reservoir, water ice was essentially in the amorphous form and the other volatiles remained trapped in the amorphous matrix ([Notesco and Bar-Nun, 2005](#)). In this context, to explain the accretion of amorphous planetesimals by the form-

ing Jupiter, [Owen et al. \(1999\)](#) proposed that either the giant planet was formed at large heliocentric distances where the temperature always favored the preservation of amorphous ice in the disk, or the protosolar nebula was much cooler at the current location of Jupiter ( $\sim 5$  AU) than predicted by current turbulent accretion disk models. In both cases, the icy material originated from the protosolar cloud and survived the formation of the protosolar nebula. This is the fraction of the icy planetesimals that vaporized when entering the envelopes of the growing Jupiter, which engendered the observed volatile enrichments. If correct, this scenario predicts that the volatile enrichments at Saturn should be in solar proportions, as seems to be the case in Jupiter, given the size of the error bars of measurements.

#### *2.4.2. Crystalline Ice Scenario*

An alternative interpretation of the volatile enrichments measured in Jupiter is the one proposed by [Gautier et al. \(2001\)](#) and subsequent papers by [Hersant et al. \(2004\)](#), [Alibert et al. \(2005\)](#) and [Mousis et al. \(2006\)](#). These authors assumed that Jupiter’s building blocks formed in the inner 30 AU of the disk, in which the gas phase has been enriched at early epochs by the vaporization of amorphous ice entering from the ISM. During the cooling of this region of the disk, water vapor crystallized and trapped the volatiles in the form of clathrates or hydrates in the 40–90 K range instead of condensing at lower temperatures. These ices then agglomerated and formed the planetesimals that were ultimately accreted by the growing Jupiter. These scenarios postulate that the amount of available crystalline water ice was large enough (typically  $\text{H}_2\text{O}/\text{H}_2 \geq 2 \times (\text{O}/\text{H})_{\odot}$ ) to trap the other volatiles in the feeding zone of Jupiter and that the disk’s temperature at which the

ices formed never decreased below  $\sim 40$  K.

Subsequent works have shown that it is possible to explain the volatile enrichments in Jupiter via the accretion and the vaporization in its envelope of icy planetesimals made from a mixture of clathrates and pure condensates (Mousis et al., 2009, 2012). These planetesimals could have formed if the initial disk’s gas phase composition was fully protosolar (including oxygen), and if the disk’s temperature decreased down to  $\sim 20$  K at their formation location. The model described in Mousis et al. (2009, 2012) is used here to show fits of the volatile enrichments measured at Jupiter and Saturn, which have been updated by using the recent protosolar abundances of Lodders et al. (2009) (see Table 3). With this model, we first computed the composition of planetesimals condensed from two extreme gas phase compositions of the nebula, namely oxidizing and reducing states. In the oxidizing state, oxygen, carbon, nitrogen are postulated to exist only in the molecular species  $\text{H}_2\text{O}$ ,  $\text{CO}$ ,  $\text{CO}_2$ ,  $\text{CH}_3\text{OH}$ ,  $\text{CH}_4$ ,  $\text{N}_2$ , and  $\text{NH}_3$ . We fixed  $\text{CO}/\text{CO}_2/\text{CH}_3\text{OH}/\text{CH}_4 = 70/10/2/1$  and  $\text{N}_2/\text{NH}_3 = 10$  in the gas phase of the disk, values usually used for the protosolar nebula (Mousis et al., 2009, 2012). In contrast, in the reducing state, C exists only in  $\text{CH}_4$  form and  $\text{N}_2/\text{NH}_3 = 0.1$  in the gas phase (Johnson et al., 2012). In both cases, P is in the form of  $\text{PH}_3$  and the volatile fraction of S is assumed to exist in the form of  $\text{H}_2\text{S}$ , with  $\text{H}_2\text{S}/\text{H}_2 = 0.5 \times (\text{S}/\text{H}_2)_\odot$ , the other fraction of S being essentially trapped in the form of troilite mineral in the solar nebula (Pasek et al., 2005). Planetesimals formed during the cooling of the nebula from these two extreme gas phase compositions are assumed to have been accreted by proto-Jupiter and proto-Saturn and devolatilized in the envelopes during their growth phases.

Once the composition of the planetesimals is defined, the adjustment of their masses accreted in the envelopes of Jupiter and Saturn allows one to determine the best fit of the observed volatile enrichments.

Figures 2 and 3 represent the fits of the enrichments observed in Jupiter’s and Saturn’s atmospheres, respectively. In the case of Jupiter, C, N, S, Ar and Kr measurements are matched by our fits, irrespective of the redox status of the protosolar nebula. Also, in both redox cases, the measured P abundance is not matched by the fits but this might be due to the difficulty of getting a reliable measurement since the mid-infrared spectrum is dominated by tropospheric ammonia. Also the measured P is predicted to be lower than its bulk abundance due to disequilibrium processes in the Jovian atmosphere (Fletcher et al., 2009a). On the other hand, Xe is almost matched by our fit in the reducing case only. The oxygen abundance is predicted to be 5.4–5.7 and 6.5–7.9 times protosolar in Jupiter in the oxidizing and reducing cases, respectively.

In the case of Saturn, our strategy was to fit the measured C enrichment. Interestingly, contrary to Jupiter, P is matched in Saturn, irrespective of the redox status of the nebula. On the other hand, the P determination is more robust in Saturn than in Jupiter because  $\text{PH}_3$  dominates the mid-infrared spectrum. However, S is not matched by our model but this might result from the lack of reliability of its determination. In addition, with enrichments predicted to be  $\sim 6$ –7 times and 11–14 times the protosolar value in the oxidizing and reducing cases, respectively, our model overestimates the amount of nitrogen present in Saturn’s atmosphere compared to observations that suggest a more moderate enrichment, in the order of  $\sim 0.5$ –4.6 times



the protosolar value. One possibility that could explain this discrepancy is that all  $\text{NH}_3$  and only a fraction of  $\text{N}_2$  would have been incorporated in Saturn’s building blocks because of the limited amount of available water favoring its efficient trapping in clathrates. The remaining fraction of  $\text{N}_2$  would have remained in the  $\text{H}_2$ -dominated gas phase of Saturn’s feeding zone as a result of the disk’s cooling down to temperatures higher than that of  $\text{N}_2$  condensation or trapping in clathrates, as proposed by [Hersant et al. \(2008\)](#). These conditions could lead to a moderate N enrichment comparable to the measured one and to a  $^{14}\text{N}/^{15}\text{N}$  ratio in the envelope very close to the Jovian value. Our model also gives predictions of O, Ar, Kr and Xe enrichments in the two redox cases. In particular, the oxygen abundance is predicted to be 14.7–18.1 and 17.5–21.5 times protosolar in the oxidizing and reducing cases, respectively. If the determination of N is confirmed at Saturn, it would appear inconsistent with the scenario proposed by [Owen et al. \(1999\)](#) because the latter predicts a uniform enrichment in volatiles in the giant planet’s envelope, which is not the case here since C/N is found to be at least  $\sim 2 \times (\text{C/N})_{\odot}$ . On the other hand, both scenarios predict the same  $^{14}\text{N}/^{15}\text{N}$  ratio at Jupiter and Saturn as the two planets accreted their nitrogen essentially from the same volatile reservoirs.

#### *2.4.3. Scenario of Supply of Refractory Carbonated Material*

[Lodders \(2004\)](#) proposed the formation of Jupiter from refractory carbonated materials, namely “tar”, placing its formation location on a “tar line” in the protosolar nebula. This scenario was used to explain the elemental abundances enrichments observed by Galileo after having normalized all the heavy elements abundances with respect to Si instead of  $\text{H}_2$ . By doing so,

Lodders (2004) found that the relative abundances of Ar, Kr, Xe and P are solar, C and possibly N are enriched, and H, He, Ne, and O are subsolar, with the Galileo H<sub>2</sub>O determination assumed to be representative of the planet's bulk O/H. In this model, Ar, Kr and Xe would have been supplied to Jupiter via direct gravitational capture of the solar nebula gas. To explain the Ar, Kr and Xe enrichments in the Jovian atmosphere, Lodders (2004) proposed that they would have been the consequence of the H<sub>2</sub> and He depletion in the envelope, which produced the metallic layer. If Saturn formed following this scenario, a useful test would be the determination of the H<sub>2</sub>O bulk abundance, which should be subsolar, as proposed by Lodders (2004) for Jupiter.

#### *2.4.4. Scenario of Disk's Gas Phase Enrichment*

To account for the enrichments in heavy noble gases observed in Jupiter's atmosphere, Guillot and Hueso (2006) proposed that Ar, Kr and Xe have condensed at  $\sim 20\text{--}30$  K onto the icy amorphous grains that settled in the cold outer part of the disk nebula midplane. These noble gases would have been released in gaseous form in the formation region of giant planets at a time when the disk would have been chemically evolved due to photoevaporation. The combination of these mechanisms would have led to a heavy noble gas enrichment relative to protosolar in the disk's gas phase from which the giant planets would have been accreted. In Guillot and Hueso (2006)'s scenario, the noble gas enrichment would have been homogeneous in the giant planets formation region. Therefore, their model predicts that the Ar, Kr and Xe enrichments in Saturn's atmosphere are similar to those observed in Jupiter, which are between  $\sim 1.8$  and 3.5 times the protosolar value (see Table 3).

These values are substantially smaller than those predicted by the model used in Sec. 2.4.2, which are in the  $\sim 4.8$ – $14.6$  times protosolar range, depending on the considered species (see Fig. 3).

### 2.5. Summary of Key Measurements

Here we provide the “recommended” measurements in Saturn’s atmosphere that would allow disentangling between i) the afore-mentioned giant planets formation scenarios and ii) the different volatile reservoirs from which the solar system bodies assembled:

- The atmospheric fraction of  $\text{He}/\text{H}_2$  with a 2% accuracy on the measurement (same accuracy as the one made by the Jamin-Mascart interferometer aboard Galileo);
- The elemental enrichments in cosmogenically abundant species C, N, S and O.  $\text{C}/\text{H}$ ,  $\text{N}/\text{H}$ ,  $\text{S}/\text{H}$  and  $\text{O}/\text{H}$  should be sampled with an accuracy better than  $\pm 10\%$  (uncertainties of the order of protosolar abundances).
- The elemental enrichments in minor species delivered by vertical mixing (e.g., P, As, Ge) from the deeper troposphere (see also Sec. 3).  $\text{P}/\text{H}$ ,  $\text{As}/\text{H}$  and  $\text{Ge}/\text{H}$  should be sampled with an accuracy better than  $\pm 10\%$  (uncertainties of the order of protosolar abundances).
- The isotopic ratios in hydrogen ( $\text{D}/\text{H}$ ), oxygen ( $^{18}\text{O}$ ,  $^{17}\text{O}$  and  $^{16}\text{O}$ ), carbon ( $^{13}\text{C}/^{12}\text{C}$ ) and nitrogen ( $^{15}\text{N}/^{14}\text{N}$ ), to determine the key reservoirs for these species (e.g., delivery as  $\text{N}_2$  or  $\text{NH}_3$  vastly alters the  $^{15}\text{N}/^{14}\text{N}$  ratio in the giant planet’s envelope).  $^{13}\text{C}/^{12}\text{C}$ ,  $^{18}\text{O}/^{16}\text{O}$  and  $^{17}\text{O}/^{16}\text{O}$

should be sampled with an accuracy better than  $\pm 1\%$ . D/H,  $^{15}\text{N}/^{14}\text{N}$  should be analyzed in the main host molecules with an accuracy of the order of  $\pm 5\%$ .

- The abundances and isotopic ratios for the chemically inert noble gases He, Ne, Xe, Kr and Ar, provide excellent tracers for the materials in the subreservoirs existing in the protosolar nebula. The isotopic ratios for He, Ne, Xe, Kr and Ar should be measured with an accuracy better than  $\pm 1\%$ .

The depth of probe penetration will determine whether it can access the well-mixed regions for key condensable volatiles. In the case of Saturn, a shallow probe penetrating down to 5–10 bar would *in principle* sample ammonia and  $\text{H}_2\text{S}$  both within and below their cloud bases, in the well-mixed regions of the atmosphere to determine the N/H and S/H ratios, in addition to noble gases and isotopic ratios. Note that the N determination could be a lower limit because ammonia is highly soluble in liquid water. Rain generated in the water cloud can provide a downward transport mechanism for ammonia, so the ammonia abundance above the water cloud could be less than the bulk abundance. Because the hypothesized water cloud is deeper than at least  $\sim 12.6$  bar in Saturn (Atreya et al., 1999), the prospect of reaching the deep O/H ratio remains unlikely if the probe would not survive beyond its design limit, unless a precise determination of the CO abundance (or any other species limited by reactions with the tropospheric water) is used to constrain  $\text{H}_2\text{O}/\text{H}_2$  (see Sec. 2.2) and/or the probe is accompanied by remote sensing experiments on a carrier spacecraft capable of probing these depths (e.g., the Juno microwave radiometer, currently en route to Jupiter). Nevertheless,

measuring elemental abundances (in particular He, noble gases and other cosmogenically-common species) and isotopic ratios using a shallow entry probe on Saturn will provide a vital comparison to Galileo’s measurements of Jupiter, and a crucial “ground-truth” for the remote sensing investigations by the Cassini spacecraft.

### **3. *In situ* Studies of Saturn’s Atmospheric Phenomena**

The giant planets are natural planetary-scale laboratories for the study of fluid dynamics without the complicating influences of terrestrial topography or ocean-atmosphere coupling. However, remote sensing only provides access to limited altitude ranges where spectral lines are formed and broadened. Furthermore, the vertical resolution of “nadir” remote sensing is fundamentally limited to the width of the contribution function (i.e., the range of altitudes contributing to the upwelling radiance at a given wavelength), which can extend over a broad range of pressures. Ground-based observatories, space telescopes and the visiting Pioneer, Voyager and Cassini missions have exploited wavelengths from the ultraviolet to the microwave in an attempt to reconstruct Saturn’s atmospheric structure in three dimensions. These studies have a limited vertical resolution and principally use visible and infrared observations in the upper troposphere (just above the condensate clouds and within the tropospheric hazes) or the mid-stratosphere near the 1 mbar level via mid-infrared emissions. Regions below the top-most clouds and in the middle/upper atmosphere are largely inaccessible to remote sensing, limiting our knowledge of the vertical variations of temperatures, densities, horizontal and vertical winds and waves, compositional profiles and cloud/haze

properties. *In situ* exploration of Saturn would not only help constrain the bulk chemical composition of this gas giant (e.g., Section 2), but it would also provide direct sampling and “ground-truth” for the myriad physical and chemical processes at work in Saturn’s atmosphere.

In the following sections we describe how an *in situ* probe, penetrating from the upper atmosphere ( $\mu$ bar pressures) into the convective weather layer to a minimum depth of 10 bar, would contribute to our knowledge of Saturn’s atmospheric structure, dynamics, composition, chemistry and cloud-forming processes. These results would be directly compared to our only other direct measurement of a giant planet, from the descent of the 339-kg Galileo probe into the atmosphere of Jupiter on December 7th 1995. The Galileo probe entered a region of unusual atmospheric dynamics near  $6.5^\circ\text{N}$ , where it is thought that the meteorology associated with planetary wave activity conspired to deplete Jupiter’s atmosphere in volatiles (e.g., [Showman and Dowling, 2000](#); [Friedson, 1999](#)), most notably preventing the probe from reaching the depth of Jupiter’s well-mixed  $\text{H}_2\text{O}$  layer after its 60-minute descent to the 22 bar level, 150 km below the visible cloud-tops. In the decade that followed, researchers have been attempting to reconcile global remote sensing of Jupiter with this single-point measurement (e.g., [Roos-Serote et al., 2000](#)). Along with the GPMS and HAD instruments, the probe carried a net flux radiometer for the thermal profile and heat budget (NFR, [Sromovsky et al., 1998](#)); a nephelometer for cloud studies (NEP, [Ragent et al., 1998](#)) and an Atmospheric Structure Instrument (ASI, [Seiff et al., 1998](#)) to measure profiles of temperature, pressure and atmospheric density. Measurements of the probe’s transmitted radio signal (driven by an ultra-stable oscillator) allowed

a reconstruction of the zonal winds with altitude (Doppler Wind Experiment, DWE, [Atkinson et al., 1998](#)), and attenuation of the probe-to-orbiter signal also provided information on the microwave opacity due to ammonia absorption ([Folkner et al., 1998](#)). Comparable *in situ* data for Saturn, in tandem with the wealth of remotely-sensed observations provided by Cassini, would enable a similar leap in our understanding of the solar system’s second giant planet. Finally, from the perspective of comparative planetology, improving our understanding of Saturn will provide a valuable new context for Galileo probe’s measurements at Jupiter, enhancing our knowledge of this unique class of planets.

### *3.1. Saturn’s Dynamics and Meteorology*

Saturn’s atmosphere stands in contrast to Jupiter, with fewer large-scale vortices and a more subdued banded structure in the visible, superimposed onto hemispheric asymmetries in temperatures, cloud cover and gaseous composition as a result of Saturn’s seasonal cycles (unlike Jupiter, Saturn has a considerable axial tilt of  $26^\circ$ ). See [West et al. \(2009\)](#), [Fouchet et al. \(2009\)](#), [Del Genio et al. \(2009\)](#) and [Nagy et al. \(2009\)](#) for detailed reviews. Despite this globally-variable atmosphere in the horizontal, a single entry probe would provide unique insights in the vertical dimension by characterising the changing environmental conditions and dynamical state as it descends from the stably-stratified middle atmosphere to the convectively-unstable troposphere. Although *in situ* probes may seem to provide one-dimensional vertical results, a horizontal dimension is also provided by reconstructing the probe trajectory during its descent, as it is buffeted by Saturn’s powerful jet streams and eddies.

### 3.1.1. Atmospheric Stability and Transition Zones

A key parameter that serves as a diagnostic of the local dynamical state of the atmosphere is the Richardson number  $Ri$ :

$$Ri = \frac{N_B^2}{\left(\frac{\partial u}{\partial z}\right)^2 + \left(\frac{\partial v}{\partial z}\right)^2} = \frac{\frac{g}{\theta} \left(\frac{\partial \theta}{\partial z}\right)}{\left(\frac{\partial u}{\partial z}\right)^2 + \left(\frac{\partial v}{\partial z}\right)^2} \quad (1)$$

where  $N_B$  is the frequency with which an air parcel would oscillate vertically due to buoyancy forces if perturbed from rest, also known as the *Brunt Väisälä* frequency;  $\theta$  is the potential temperature and  $\frac{\partial \theta}{\partial z}$  the static stability;  $g$  is the gravitational acceleration and  $u$  and  $v$  the zonal and meridional velocities, respectively. An entry probe can measure continuous profiles of these parameters as a function of altitude, enabling a study of stability and instability regimes as a function of depth. Temperatures and densities in the upper atmosphere can be determined via the deceleration caused by atmospheric drag, connecting the high temperature thermosphere at nanobar pressures to the middle atmosphere at microbar and millibar pressures (e.g., [Yelle and Miller, 2004](#)). An atmospheric structure instrument would measure atmospheric pressures and temperatures throughout the descent to the clouds, and from these infer atmospheric stability and densities (provided the mean molecular weight is determined by another instrument; [Seiff et al., 1998](#); [Magalhães et al., 2002](#)). Upper atmospheric densities would be deduced from measured accelerations and from area and drag coefficients<sup>2</sup>. The probe will sample both the radiatively-cooled upper atmosphere and also the

---

<sup>2</sup>Note that ablation sensors on the entry probe are needed to get the time-profile of Thermal Protection System (TPS) mass loss and change in area during entry.



convectively driven troposphere, precisely constraining the static stability, radiative-convective boundary (i.e., how far down does sunlight penetrate?) and the levels of the tropopause, stratopause, mesopause and homopause. Thermal structure measurements of Saturn would be directly compared to those on Jupiter to understand the energetic balance between solar heating, thermal cooling, latent heat release, wave heating and internal energy for driving the complex dynamics of all the different atmospheric layers on the giant planets, and how this balance differs as a function of distance from the Sun.

### *3.1.2. Wave Activity*

Perturbations of the temperature structure due to vertical propagation of gravity waves are expected to be common features of the stably stratified middle atmospheres either on terrestrial planets or gas giants. Wave activity is thought to be a key coupling mechanism between the convective troposphere (e.g., gravity waves and Rossby/planetary waves generated by rising plumes and vortices) and the stable middle/upper atmosphere, being responsible for transporting energy and momentum through the atmosphere and for phenomenon like the Quasi-Biennial Oscillation on Earth ([Baldwin et al., 2001](#)), which is thought to have counterparts on Jupiter and Saturn ([Fouchet et al., 2008](#)). Waves are a useful diagnostic of the background state of the atmosphere, as their propagation relies on certain critical conditions (e.g., the static stability and vertical shears on zonal winds, which cannot be revealed by remote sensing alone). Energy and momentum transfer via waves serve as a source of both heating and cooling for the hot thermospheres, whose temperatures far exceed the expectations from solar heating alone, although

the precise origins of the heating source has never been satisfactorily identified (e.g., [Hickey et al., 2000](#); [Nagy et al., 2009](#)). The periodicity of gravity waves measured by the Galileo probe on Jupiter permits the reconstruction of the zonal wind profile from the lower thermosphere to the upper troposphere ([Watkins and Cho, 2013](#)), permitting identification of the homopause (where molecular and eddy diffusion become comparable and gravity waves break to deposit their energy), above which the atmosphere separates into layers of different molecular species. Understanding the propagation, periodicity and sources of wave activity on Saturn will reveal the properties of the background medium and the coupling of the “weather layer” to the middle atmosphere especially on how zonal and meridional circulations are forced by eddy-mean flow interactions, and facilitate direct comparison with Jupiter.

### *3.1.3. Profiling Atmospheric Winds*

*In situ* exploration would tackle one of the most enduring mysteries for the giant planets - what powers and maintains the zonal winds responsible for the planetary banding, how deep do those winds penetrate into the troposphere, and what are the wind strengths in the middle atmosphere? Remote sensing of temperature contrasts (and hence wind shears via thermal wind relationships), or inferences from the properties of atmospheric plumes at the cloud-tops (e.g., [Sánchez-Lavega et al., 2008](#)) cannot directly address this question. Remotely observed cloud motions are often ambiguous due to uncertainties in the cloud location; the clouds themselves may be imperfect tracers of the winds; and vertical temperature profiles (and hence wind shears) are degenerate with the atmospheric composition. *In situ* measurements of the vertical variation of winds and temperatures should resolve

these ambiguities. The Galileo probe’s DWE reported that jovian winds were at a minimum at the cloud tops (where most of our understanding of zonal winds and eddy-momentum fluxes originate from), and increased both above (Watkins and Cho, 2013) and below (Atkinson et al., 1998) this level. In the deep atmosphere, DWE demonstrated that Jupiter’s winds increased to a depth of around 5 bars, and then remained roughly constant to the maximum probe depth of around 22 bars. Similar measurements on Saturn could sample the transition region between two different circulation regimes - an upper tropospheric region where eddies cause friction to decelerate the zonal jets and air rises in cloudy zones, and a deeper tropospheric region where the circulation is reversed and eddy pumping is essential to maintain the jets and air rises in the warmer belts (e.g., Del Genio et al., 2009; Fletcher et al., 2011). A single entry probe would potentially sample both regimes, and reconciling these two views of tropospheric circulation on Saturn would have implications for all of the giants. Finally, direct measurements of winds in the middle atmosphere would establish the reliability of extrapolations from the jets in the cloud tops to the stratosphere in determining the general circulations of planetary stratospheres.

### *3.2. Saturn’s Clouds and Composition*

In Section 2 we discussed the need for reliable measurements of bulk elemental enrichments and isotopic ratios to study the formation and evolution of Saturn. Vertical profiles of atmospheric composition (both molecular and particulate) are essential to understanding the chemical, condensation and disequilibrium processes at work, in addition to the deposition of material from outside of the planet’s atmosphere. The Galileo probe compositional

and cloud measurements revealed an unexpectedly dry region of the jovian troposphere, depleted in clouds and volatiles (Atreya et al., 1999), which was consistent with ground-based observations of the probe entry into a warm cyclonic region (e.g., Orton et al., 1998). For this reason, the compositional profiles measured by Galileo are not thought to be globally representative of Jupiter’s atmosphere, leading to a desire for multiple entry probes for different latitudes and longitudes in future missions. Nevertheless, a single probe is essential for a more complete understanding of this class of giant planets, to enhance our knowledge of Saturn and to provide a context for improved interpretation of the Galileo probe’s sampling of Jupiter’s unusual meteorology.

### *3.2.1. Clouds and hazes*

A poor understanding of cloud and haze formation in planetary atmospheres of our solar system may be the key parameter limiting our ability to interpret spectra of extrasolar planets and brown dwarfs (e.g., Marley et al., 2013). Although equilibrium cloud condensation models (ECCMs, Weidenschilling and Lewis, 1973) combined with the sedimentation of condensates to form layers, have proven successful in explaining the broad characteristics of the planets (methane ice clouds on ice giants, ammonia ice clouds on gas giants), they remain too simplistic to reproduce the precise location, extent and microphysics of the observed cloud decks. The Galileo probe results defied expectations of equilibrium condensation by revealing clouds bases at 0.5, 1.3 and 1.6 bar, plus tenuous structure from 2.4-3.6 bar and no evidence for a deep water cloud (Atreya et al., 1999; West et al., 2004). Ammonia ice on Jupiter has only been spectroscopically identified in regions of powerful

convective updrafts (e.g., [Baines et al., 2002](#); [Reuter et al., 2007](#)), and water ice has been detected in Voyager far-infrared spectroscopy ([Simon-Miller et al., 2000](#)). The spectral signature of pure ammonia ice is likely obscured by a coating or mixing with other products, such as photolytically produced hydrocarbons, hydrazine or diphosphine (e.g., [Sromovsky and Fry, 2010](#); [West et al., 2004](#)). The spectral properties of these mixtures are poorly known, rendering cloud remote sensing highly ambiguous. Furthermore, Saturn's upper troposphere appears dominated by a ubiquitous haze whose composition has never been determined and is potentially unrelated to condensed volatiles (although diphosphine,  $\text{P}_2\text{H}_4$ , a product of the UV destruction of phosphine, remains an intriguing possibility). An ECCM applied to Saturn with a  $5\times$  enhancement of heavy elements over solar abundances predicts  $\text{NH}_3$  condensation at 1.8 bar,  $\text{NH}_4\text{SH}$  near 4 bar and an aqueous ammonia cloud (merging with a water ice cloud) near 20 bar ([Atreya et al., 1999](#)). However, ammonia and water ice signatures have been identified only recently, in the powerful updrafts associated with a powerful springtime storm in 2010–2011 ([Sromovsky et al., 2013](#)).

The only way to resolve these questions is by *in situ* sampling of the clouds and hazes formed in a planet's atmosphere, using instruments designed to measure the particle optical properties, size distributions, number and mass densities, optical depth and vertical distribution. Combined with the vertical profiles of condensable volatiles (e.g.,  $\text{NH}_3$ ,  $\text{H}_2\text{S}$  and  $\text{H}_2\text{O}$  on Saturn) and photochemically-produced species (hydrocarbons, hydrazine  $\text{N}_2\text{H}_4$ , diphosphine), this would give an estimate of the composition of Saturn's condensation clouds and upper atmospheric hazes for the first time. Sat-

urn's atmosphere provides the most accessible cloud decks for this study after Jupiter (condensates of  $\text{NH}_3$  and  $\text{H}_2\text{O}$  are locked away at considerably higher pressures on the ice giants); the most useful comparison to remote sensing data (e.g., from Cassini); and the most similar composition to Jupiter for a full understanding of gas giant clouds.

### *3.2.2. Atmospheric Chemistry and Mixing*

Gaseous species can be removed from the gas phase by condensation; modified by vertical mixing and photolysis; and deposited from exogenic sources (icy rings, satellites, interplanetary dust, comets, etc.), causing abundance profiles to vary with altitude and season. Indeed, all the giant planets exhibit a rich chemistry due to the UV photolysis of key atmospheric species. Their stratospheres are dominated by the hydrocarbon products of methane photolysis (e.g., [Moses et al., 2005](#)), which descend into the troposphere to be recycled by thermochemical conversion. On Jupiter, the Galileo probe was able to measure hydrocarbon species in the 8–12 bar region, although the balance of ethane (expected to be the most abundant hydrocarbon after methane) to ethylene, propene, acetylene and propane led to suspicions that the hydrocarbon detections were instrumental rather than of atmospheric origin ([Wong, 2009](#)). Stratospheric measurements of hydrocarbons in their production region were not performed, but would be possible on Saturn with a probe. Saturn's troposphere features saturated volatiles in trace amounts above the cloud tops, but only ammonia gas is abundant enough for remote detection.  $\text{H}_2\text{S}$  and  $\text{H}_2\text{O}$  profiles above the condensation clouds have never been measured. In addition to the volatiles, Saturn's troposphere features a host of disequilibrium species, most notably phosphine, dredged up from

the deeper, warmer interior by vigorous atmospheric mixing (e.g., [Fletcher et al., 2009a](#)). The abundance of  $\text{PH}_3$  measured in the upper troposphere is thought to represent the abundance at its thermochemical quench level, where the vertical diffusion timescale is shorter than the thermochemical kinetics timescale. Measurements of additional trace species in the troposphere ( $\text{GeH}_4$ ,  $\text{AsH}_3$ ,  $\text{CO}$ ) provide constraints on the strength of atmospheric mixing from deeper, warmer levels below the clouds.  $\text{CO}$  is of particular interest because it could be used as a probe of the deep O/H ratio of Saturn (see [Section 2](#)).

Detection of trace chemical species ( $\text{HCN}$ ,  $\text{HCP}$ ,  $\text{CS}$ , methanol, formaldehyde) and hydrogen halides ( $\text{HCl}$ ,  $\text{HBr}$ ,  $\text{HF}$  and  $\text{HI}$ , e.g., [Teanby et al., 2006](#); [Fletcher et al., 2012](#)) would reveal coupled chemistry due to lightning activity or shock chemistry due to planetary impacts. In addition, the presence of oxygenated species in the upper stratosphere ( $\text{CO}$ ,  $\text{CO}_2$ ,  $\text{H}_2\text{O}$ ) reveal the strength of exogenic influx of materials (comets, interplanetary dust, e.g., [Feuchtgruber et al., 1997](#); [Cavalié et al., 2010](#)) into the upper atmosphere of Saturn. Sensitive mass spectrometry of these species, combined with probe measurements of atmospheric temperatures and haze properties, could reveal the processes governing the soup of atmospheric constituents on the giant planets. Once again, Saturn's trace species are expected to be the most accessible of the solar system giant after Jupiter, as volatiles and disequilibrium species (e.g.,  $\text{PH}_3$  and  $\text{NH}_3$ ) have so far eluded remote detection on the ice giants.

### 3.3. Summary of Key Atmospheric Measurements

A single entry probe would reveal new insights into the vertical structures of temperatures, density, chemical composition and clouds during descent through a number of different atmospheric regions, from the stable upper/middle atmosphere to the convective troposphere. It would directly sample the condensation cloud decks and ubiquitous hazes whose composition, altitude and structure remain ambiguous due to the inherent difficulties with remote sensing. Furthermore, it would show how Saturn's atmosphere flows at a variety of different depths above, within and below the condensate clouds. Key measurements required to address the science described in this section include:

- Continuous measurements of atmospheric temperature and pressure throughout the descent to study (i) stability regimes as a function of depth through transition zones (e.g., radiative-convective boundary); (ii) atmospheric drag and accelerations; and (iii) the influence of wave perturbations and cloud formation on the vertical temperature profile;
- Determination of the vertical variation of horizontal winds using Doppler measurements of the probe's carrier frequency (driven by an ultra-stable oscillator) during the descent. This includes a study of the depth of the zonal wind fields, as well as the first measurements of middle atmospheric winds;
- Vertical profiling of a host of atmospheric species via mass spectrometry, including atmospheric volatiles (water,  $\text{H}_2\text{S}$  and  $\text{NH}_3$  in their saturated and sub-cloud regions); disequilibrium species (e.g.,  $\text{PH}_3$ ,  $\text{AsH}_3$ ,



GeH<sub>4</sub>, CO) dredged from the deeper atmosphere; photochemical species (e.g., hydrocarbons and HCN in the troposphere and stratosphere; hydrazine and diphosphine in the upper troposphere) and exogenic inputs (e.g., oxygenated species in the upper atmosphere);

- Measurements of the vertical structure and properties of Saturn’s cloud and haze layers; including determinations of the particle optical properties, size distributions, number and mass densities, opacity, shapes and, potentially, their composition.

With a single entry probe, the selected entry site must be carefully studied. Saturn’s equatorial zone is one potential site for a single entry probe because of its meteorological activity that combines: the emergence of large-scale storms ([Sanchez-Lavega et al., 1991](#)); vertical wind shears in the troposphere ([García-Melendo et al., 2011](#)); upwelling enhancing volatiles and disequilibrium species ([Fletcher et al., 2009a, 2011](#)); and a global stratospheric oscillation of the thermal field ([Fouchet et al., 2008](#); [Orton et al., 2008](#); [Guerlet et al., 2011](#)). Additionally, the strength of its equatorial eastward jet (peak velocities up to 500 m/s) poses one of the theoretical challenges to the understanding of the dynamics of fluid giant planets. Furthermore, a descent probe into Saturn’s equatorial region could further constrain the influx of H<sub>2</sub>O originating from the Enceladus torus ([Hartogh et al., 2011](#)). However, it remains an open question as to how representative the equatorial region would be of Saturn’s global dynamics. Short of multiple entry probes targeted at different regions of upwelling and subsidence, near to narrow prograde jets or broader retrograde jets, a mid-latitude atmospheric region might be a more representative sample.

#### 4. Mission Architectures

The primary science objectives described in Sec. 2 and 3 may be addressed by an atmospheric entry probe that would descend under parachute, and start to perform *in situ* measurements in the stratosphere to help characterize the location and properties of the tropopause, and continue into the troposphere to pressures of at least 10 bars. All of the science objectives, except for the abundance of oxygen which may be only addressed partially, can be achieved by reaching 10 bars. Previous studies have shown that depths beyond 10 bars become increasingly more difficult to achieve for several technology reasons; for example: i) the descent time, hence the relay duration, would increase and make the relay geometry more challenging; ii) the technology for the probe may change at pressures greater than 10 bars; iii) the opacity of the atmosphere to radio-frequencies increases with depth and may make the communication link even more challenging at higher pressures. Future studies would be needed to conduct a careful assessment of the trade-offs between science return and the added complexity of a probe that could operate at pressures greater than 10 bars. Accelerometry measurements may also be performed during the entry phase in the higher part of the stratosphere to probe the upper layers of the atmosphere prior to starting *in situ* measurements under parachute.

A carrier spacecraft would be required to deliver the probe to the desired atmospheric entry point at Saturn. We have identified three possible mission configurations:

- **Configuration 1: Probe + Carrier.** The probe would detach from the carrier spacecraft prior to probe entry. The carrier would follow

the probe path and be destroyed during atmospheric entry, but may be capable of performing pre-entry science. The carrier would not be used as a radio relay to transmit the probe data to Earth. The probe would transmit its data to the ground system via a direct-to-Earth (DTE) RF link;

- **Configuration 2: Probe + Carrier/Relay.** The probe would detach from the carrier several months prior to probe entry. Subsequent to probe release, the carrier trajectory would be deflected to prepare for over-flight phasing of the probe descent location for both probe data relay as well as performing approach and flyby science;
- **Configuration 3: Probe + Orbiter.** This configuration would be similar to the Galileo Orbiter/Probe mission. The probe would detach from the orbiter several months prior to probe entry. As for Configuration 2, subsequent to probe release, the orbiter trajectory would be deflected to prepare for over-flight phasing of the probe descent location. After probe relay during over-flight, the orbiter would be placed in orbit around Saturn and continue to perform orbital science.

Configuration 1 would allow the carrier to perform months of approach science and *in situ* pre-entry science. In this architecture, the probe data transmission would rely solely on a Direct-to-Earth probe telecommunications link. In addition to being used as the probe relay data following completion of the probe mission, Configuration 2 would possibly also provide the capability to perform months of approach science, but in addition flyby science (for a few days). This configuration would allow many retransmis-

sions of the probe data for redundancy. Configuration 3 would clearly be the most capable, but most costly configuration. Trade-off studies will need to be carried out to assess whether the supporting remote sensing observations may be achievable during the approach phase and a single flyby or from an orbiter. Any of the carrier options could provide context observations but an orbiter could bring more science return in addition to supporting the probe science. The only requirement is that those data be downlinked to Earth while the spacecraft is still operating. For example, useful observations from a Configuration 1 carrier could be made several hours before probe entry, and downlink could be accomplished in the intervening time. Finally, it may be worth studying if the emerging solar-sail propulsion technology ([Janhunen et al., 2014](#)) can be considered for this option.

#### *4.1. Atmospheric Entry Probe*

An atmospheric entry probe at Saturn would in many respects resemble the Jupiter Galileo probe. The concept was put forward for Saturn in the KRONOS mission proposal (Marty et al. 2009). Giant Planet probe concept studies have been studied by ESA in 2010<sup>3</sup>. As an example, the KRONOS probe had a mass of  $\sim 337$ kg, with a 220kg deceleration module (aeroshell, thermal protection system, parachutes and separation hardware) and a 117kg descent module, including the probe structure, science instruments, and subsystems. It is anticipated that the probe architecture for this mission would be battery powered and accommodate either a DTE link or a data relay to the carrier or the orbiter. Trades would be done to assess the complexity

---

<sup>3</sup><http://sci.esa.int/sre-fp/47568-pep-assessment-study-internal-final-presentation/>

(and cost) of probe and telecomm link design as a function of operational depth in the atmosphere. A representative payload for the Saturn probe that would allow addressing the science objectives identified in Sec. 2 and 3 is shown in Table 5.

#### *4.2. Carrier or Orbiter*

Alternative architectures for the carrier (Configuration 1 or 2) or the orbiter (Configuration 3) would be considered, taking into account, if possible and if technologically and programmatically sound, the heritage for outer planet/deep space missions within either ESA or NASA. As an example, the carrier or the orbiter may benefit from subsystems developed by either ESA or NASA for previous outer planet missions (for example ESA/JUICE or NASA/JUNO, or possibly NASA/ESA Cassini-Huygens).

#### *4.3. Power Generation*

It would be worth studying whether the proposed mission architectures could be solely designed on batteries and solar power. It would require qualification of the low-intensity low-temperature (LILT) solar array cells for 9.5 AU conditions. The probe would be powered with primary batteries as were the Galileo and Huygens probes. In all three configurations, the carrier (configuration 1 and 2) or the orbiter (configuration 3) would be equipped with a combination of solar panels, secondary batteries and possibly a set of primary batteries for phases that require a high power demand, for example during the probe entry phase. Nuclear power would be considered for the carrier or the orbiter only if available solar power technology would be found to be unfeasible.

#### *4.4. Interplanetary Trajectory and Entry Zone of the Probe*

Many trajectory options have been identified, using both direct and gravity-assisted transfers to Saturn, and more will be identified in subsequent studies. Trajectory selection will be based on the selected carrier option, launch vehicle capabilities, and available probe thermal protection capability. The interplanetary trajectory and the probe entry location are inseparably linked. Saturn's extensive ring system presents a severe collision hazard to an inbound probe. For various declinations of the spacecraft's approach asymptote, some latitudes will be inaccessible because the trajectories to deliver to those latitudes would impact the rings. Also, although it is possible to adjust the inclination of the approach orbit for purposes of accessing a desired latitude, this approach can greatly increase the atmosphere-relative entry speeds, possibly driving the mission to an expensive heat shield material technology development. During the studies, the issues of probe entry locations, approach and entry trajectories, and probe technologies must be treated together. Due to Saturn's large obliquity and the characteristics of reasonable Earth-to-Saturn transfer trajectories, the best combinations change with time. Concerning the probe entry zone, both equatorial and mid-latitude regions may be a representative location from the scientific point of view (see a discussion in Sec. 3.3). Volatile-depleted regions are probably located at the cyclones in both poles and may also be located at the so-called "storm-alley" (region of low static stability able to develop updrafts and downdrafts). More generally, the peaks of westward jets can be unstable based on the stability of the wind system and eastward jets (particularly the anticyclonic branch of eastward jets) might be good locations to retrieve the deep values of volatiles at higher

levels in the atmosphere (Read et al., 2009). In any case, there are several potential entry points and a dedicated study will be needed to assess their relative merit.

#### *4.5. International Collaboration*

In this paper, we only consider ESA/Europe and NASA/USA collaborations but collaborations with other international partners may be envisaged. One of the key probe technologies for a Saturn probe that would be new for European industry, is the heat shield material. Recent NASA studies concerning entry system performance requirements versus thermal protection system capability for a Saturn entry probe have been completed (Ellerby et al., 2013). This example is used to illustrate that, should Europe be willing to lead the probe development (as was so successfully done with Huygens), careful consideration of trade-offs would have to be made for either development of this new technology within Europe or for establishing an international collaboration with NASA. International collaboration may also be considered for other mission elements, including the carrier (and of course the orbiter if configuration 3 would be further studied), navigation, the probe data relay, the ground segment, and science payload. All three configurations would be achievable through different schemes of collaboration between Europe and the USA. As an example, configurations 1 and 2 may take the form of a combined ESA/Class-M and a NASA Discovery or New Frontiers collaboration, if such a scheme were to become programmatically feasible as it is currently not the case. Configuration 3 would be achievable through a collaboration that would involve an ESA/Class M and a NASA/Flagship mission. We do not put forward an ESA/Class L mission at this stage as the current ESA

Cosmic Vision plan would not allow a new Class-L mission before the late 30's/early 40's.

## 5. Characteristics of a Possible Probe Model Payload

The scientific requirements discussed above can be addressed with a suite of scientific instruments, which are given in Table 5 and discussed in the following. Note that this list of instruments should not be considered as a unique payload complement but as guideline for some of the instruments we might wish to see on board. For example, an alternative to both the nephelometer and net flux radiometer described below are elements of the Huygens Descent Imager/Spectral Radiometer (DISR) (Tomasko et al., 2002) that used the sun as a source. Ultimately, the payload of the Saturn probe would be the subject of detailed mass, power and design trades, but should seek to address the majority of the scientific goals outlined in Sec. 2 and 3.

### 5.1. Mass Spectrometry

The chemical and isotopic composition of Saturn's atmosphere, and its variability, will be measured by mass spectrometry. The gas analysis systems for a Saturn Probe may benefit from a high heritage from instrumentation already flown and having provided atmospheric composition and isotope investigations. The scientific objective for the mass spectrometric investigation regarding Saturn's formation and the origin of the solar system are *in situ* measurements of the chemical composition and isotope abundances in the atmosphere, such as H, C, N, S, P, Ge, As, noble gases He, Ne, Ar, Kr, and Xe, and the isotopes D/H,  $^{13}\text{C}/^{12}\text{C}$ ,  $^{15}\text{N}/^{14}\text{N}$ ,  $^3\text{He}/^4\text{He}$ ,  $^{20}\text{Ne}/^{22}\text{Ne}$ ,  $^{38}\text{Ar}/^{36}\text{Ar}$ ,  $^{36}\text{Ar}/^{40}\text{Ar}$ , and those of Kr and Xe.



At Jupiter, the Galileo Probe Mass Spectrometer (GPMS) experiment (Niemann et al., 1992) was used, which was designed to measure the chemical and isotopic composition of Jupiter’s atmosphere in the pressure range from 0.15 to 20 bar by *in situ* sampling of the ambient atmospheric gas. The GPMS consisted of a gas sampling system that was connected to a quadrupole mass spectrometer. The gas sampling system also had two sample enrichment cells, allowing for enrichments of hydrocarbons by a factor 100 to 500, and one noble gas analysis cell with an enrichment factor of about 10. Low abundance noble gases can be measured by using an enrichment cell, as used on the Galileo mission (Niemann et al., 1996). From GPMS measurements the Jupiter He/H<sub>2</sub> ratio was determined as  $0.156 \pm 0.006$ . To improve the accuracy of the measurement of the He/H<sub>2</sub> ratio and isotopic ratios by mass spectrometry the use of reference gases will be necessary. On the Rosetta mission the ROSINA experiment carries for each mass spectrometer a gas calibration unit (Balsiger et al., 2007). Similarly, the SAM experiment on the Curiosity rover can use either a gas sample from its on-board calibration cell or utilize one of the six individual metal calibration cups on the sample manipulation system (Mahaffy et al., 2012).

A major consideration for the mass spectrometric analysis is how to distinguish between different molecular species with the same nominal mass, e.g. N<sub>2</sub> and CO, from the complex mixture of gases in Saturn’s atmosphere. There are two approaches available, one is high resolution mass spectrometry with sufficient mass resolution to resolve the isobaric interferences, and the other is chemical pre-separation of the sample followed by low resolution mass spectrometry.

### 5.1.1. High Resolution Mass Spectrometry

Probably the first composition experiment with high resolution mass spectrometry is the ROSINA experiment on the Rosetta mission (Balsiger et al., 2007) which has a Reflectron-Time-of-Flight (RTOF) instrument with a mass resolution of about  $m/\Delta m = 5,000$  at 50% peak height (Scherer et al., 2006), Double-Focussing Mass Spectrometer (DFMS) with a mass resolution of about  $m/\Delta m = 9,000$  at 50% peak height, and a pressure gauge. Determination of isotope ratios at the 1% accuracy level has been accomplished during the cruise phase. A time-of-flight instrument with even more mass resolution has been developed for possible application in Titan's atmosphere, which uses a multi-pass time-of flight configuration (Waite et al., 2012). Typical mass resolutions are  $m/\Delta m = 13,500$  at 50% peak height and 8,500 at 20% peak height. In bunch mode the mass resolution can be increased to 59,000 (at 50% peak height). Again, determination of isotope ratios at the 1% accuracy level has been accomplished. An alternative multi-pass time-of-flight instrument has been developed by Okumura et al. (2004), which uses electric sectors instead of ion mirrors for time and space focussing. Mass resolutions up to  $m/\Delta m = 350,000$  have been reported (Toyoda et al., 2003).

Recently, a new type of mass spectrometer, the Orbitrap mass spectrometer, was introduced (Makarov, 2000; Hu et al., 2005), which uses ion confinement in a harmonic electrostatic potential. The Orbitrap mass spectrometer is a Fourier-Transform type mass spectrometer, and it allows for very high mass resolutions in a compact package. For example, using an Orbitrap mass spectrometer for laboratory studies of chemical processes in Titan's atmo-

sphere, mass resolutions of  $m/\Delta m = 100,000$  have been accomplished up to  $m/z = 400$  (Hörst et al., 2012), and  $m/\Delta m = 190,000$  at 50% peak height and  $m/z = 56$  in a prototype instrument for the JUICE mission (Briois et al., 2013).

### 5.1.2. Low Resolution Mass Spectrometry with Chemical Pre-Processing

The alternative approach to high resolution mass spectrometry, which was used successfully on many missions so far, is to use a simpler low resolution mass spectrometer together with a chemical processing of the sample to separate or eliminate isobaric interferences. One established way is to use chromatographic columns with dedicated chemical specificity for a separation of chemical substances before mass spectrometric analysis. The Gas-Chromatograph Mass Spectrometer (GCMS) of the Huygens Probe is a good example of such an instrument (Niemann et al., 2002, 2005, 2010). The Huygens Probe GCMS has three chromatographic columns, one column for separation of CO and N<sub>2</sub> and other stable gases, the second column for separation of nitriles and other organics with up to three carbon atoms, and the third column for the separation of C<sub>3</sub> through C<sub>8</sub> saturated and unsaturated hydrocarbons and nitriles of up to C<sub>4</sub>. The GCMS was also equipped with a chemical scrubber cell for noble gas analysis and a sample enrichment cell for selective measurement of high boiling point carbon containing constituents. A quadrupole mass spectrometer was used for mass analysis with a mass range from 2 to 141 amu, which is able to measure isotope ratios with an accuracy of 1%. Newer examples of GCMS instrumentation are the Ptolemy instrument on the Rosetta lander for the measurement of stable isotopes of key elements (Wright et al., 2007), which uses an ion trap mass spectrometer,

the COSAC instrument also on the Rosetta lander for the characterization of surface and sub-surface samples (Goesmann et al., 2007), which uses a time-of-flight mass spectrometer, and the SAM experiment on the Curiosity rover (Mahaffy et al., 2012), which uses a classical quadrupole mass spectrometer.

### 5.1.3. Summary of Mass Spectrometry

So far in most missions the chemical pre-separation was the technique used to avoid isobaric interferences in the mass spectra, with the exception of the mass spectrometer experiment ROSINA on the Rosetta orbiter. Chemical pre-separation works well, but by choosing chromatographic columns with a certain chemical specificity one makes a pre-selection of the species to be investigated in detail. This possibly is a limitation when exploring an object where little is known. Also, gas chromatographic systems with several columns are rather complex systems, both to build and to operate (see the SAM instrument as a state-of-the art example of this technique (Mahaffy et al., 2012)).

In recent years there has been a significant development of compact mass spectrometers that offer high mass resolution, as outlined above, and these developments are still ongoing. Thus, solving the problem of isobaric interferences in the mass spectra by mass resolution can be addressed by mass spectrometry alone and one should seriously consider using high resolution mass spectrometry for a future mission to probe Saturn's atmosphere. After all, no *a priori* knowledge of the chemical composition has to be assumed. In addition, with modern time-of-flight mass spectrometers mass ranges beyond 1000 amu are not a problem at all, which would have been useful to investigate Titan's atmosphere. Nevertheless, some chemical pre-selection may still

be considered even for high resolution mass spectrometry. For example, the cryotrapping technique, which has a long history in the laboratory, would enable detection of noble gases at abundances as low as 0.02 ppb ([Waite et al., 2012](#)).

#### 5.1.4. Tunable Laser System

A Tunable Laser Spectrometer (TLS) ([Durry et al., 2002](#)) can be employed as part of a GC system to measure the isotopic ratios to a high accuracy of specific molecules, e.g. H<sub>2</sub>O, NH<sub>3</sub>, CH<sub>4</sub>, CO<sub>2</sub> and others. TLS employs ultra-high spectral resolution (0.0005 cm<sup>-1</sup>) tunable laser absorption spectroscopy in the near infra-red (IR) to mid-IR spectral region. TLS is a direct non-invasive, simple technique that for small mass and volume can produce remarkable sensitivities at the sub-ppb level for gas detection. Species abundances can be measured with accuracies of a few %, and isotope determinations are with about 0.1% accuracy. With a TLS system one can derive the isotopic ratios of D/H, <sup>18</sup>O/<sup>16</sup>O, <sup>13</sup>C/<sup>12</sup>C, <sup>18</sup>O/<sup>16</sup>O, and <sup>17</sup>O/<sup>16</sup>O.

For example, TLS was developed for application in the Mars atmosphere ([Le Barbu et al., 2004](#)), within the ExoMars mission; a recent implementation of a TLS system was for the Phobos Grunt mission ([Durry et al., 2010](#)), and is on the SAM instrument ([Webster and Mahaffy, 2011](#)), which was used to measure the isotopic ratios of D/H and of <sup>18</sup>O/<sup>16</sup>O in water and <sup>13</sup>C/<sup>12</sup>C, <sup>18</sup>O/<sup>16</sup>O, <sup>17</sup>O/<sup>16</sup>O, and <sup>13</sup>C<sup>18</sup>O/<sup>12</sup>C<sup>16</sup>O in carbon dioxide in the Martian atmosphere ([Webster et al., 2013](#)).

### 5.2. Helium Abundance Detector

The Helium Abundance Detector (HAD), as it was used on the Galileo mission (von Zahn and Hunten, 1992), basically measures the refractive index of the atmosphere in the pressure range of 2–10 bar. The refractive index is a function of the composition of the sampled gas, and since the jovian atmosphere consists of mostly of H<sub>2</sub> and He, to more than 99.5%, the refractive index is a direct measure of the He/H<sub>2</sub> ratio. The refractive index can be measured by any two-beam interferometer, where one beam passes through a reference gas and the other beam through atmospheric gas. The difference in the optical path gives the difference in refractive index between the reference and atmospheric gas. For the Galileo mission a Jamin-Mascart interferometer was used, because of its simple and compact design, with an expected accuracy of the He/H<sub>2</sub> ratio of  $\pm 0.0015$ . The accomplished measurement of the He mole fraction gave  $0.1350 \pm 0.0027$  (von Zahn et al., 1998), with a somewhat lower accuracy when expected, but still better than is possible by a mass spectrometric measurement.

### 5.3. Atmospheric Structure Instrument

The key *in situ* measurements by an Atmospheric Structure Instrument (ASI) would be the accelerometry during the probe entry phase and pressure, temperature and mean molecular weight during descent. The atmospheric density is derived from these measurements. There is strong heritage from the Huygens ASI experiment (HASI) of the Cassini-Huygens mission (Fulchignoni et al., 2002). Furthermore, these types of sensors have been selected for NASA's Mars Science Laboratory (MSL) and are part of the meteorological package of ESA's Exomars mission. *In situ* atmospheric structure

measurements are essential for the investigation of the planet's atmospheric structure and dynamics. The estimation of the temperature lapse rate can be used to identify the presence of condensation and possible clouds, to distinguish between saturated and unsaturated, stable and conditionally stable regions. The variations in the density, pressure and temperature profiles provide information on the atmospheric stability and stratification, on the presence of winds, thermal tides, waves and turbulence in the atmosphere. A typical Atmospheric Structure Instrument would consist of three primary sensor packages: a three-axis accelerometer, a pressure profile instrument and temperature sensors. It would start to operate right at the beginning of the entry phase of the probe, sensing the atmospheric drag experienced during entry. Direct pressure and temperature measurement could be performed by the sensors having access to the atmospheric flow from the earliest portion of the descent until the end of the probe mission at approximately 10 bar.

ASI data will also contribute to the analysis of the atmospheric composition, since it will monitor the acceleration experienced by the probe during the whole descent phase. ASI will provide the unique direct measurements of pressure and temperature through sensors having access to the atmospheric flow.

### *5.3.1. Accelerometers*

The accelerator package, a 3-axis accelerometer, should be placed as close as possible to the center of mass of the entry probe. Like on Huygens, the main sensor could be a highly sensitive servo accelerometer aligned along the vertical axis of the Probe, with a resolution of 1 to  $10 \times 10^{-5} \text{ m s}^{-2}$  with an accuracy of 1%. Probe acceleration can be measured in the range of

0–2000 m s<sup>-2</sup> (Zarnecki et al., 2004). Assuming the HASI accelerometer performance at Titan, a noise level of about  $3 \times 10^{-8}$  m s<sup>-2</sup> is expected. The exact performance achievable, in terms of the accuracy of the derived atmospheric density, will also depend on the probe ballistic coefficients, entry speed and drag coefficient.

### 5.3.2. *Temperature sensors*

As in the Huygens Probe, the temperature sensors will use platinum resistance thermometers. These are exposed to the atmospheric flow and are effectively thermally isolated from the support structure. The principle of measurement is based on the variation of the resistance of the metallic wire with temperature. TEM has been designed to have a good thermal coupling between the sensor and the atmosphere and to achieve high accuracy and resolution. Over the temperature range of 60–330 K these sensors maintain an accuracy of 0.1 K with a resolution of 0.02 K.

### 5.3.3. *Pressure Profile Instrument*

The Pressure Profile Instrument (PTI) will measure the pressure during the entire descent with an accuracy of 1% and a resolution of  $10^{-6}$  bar. The atmospheric flow is conveyed through a Kiel probe inside the probe where the transducers and related electronic are located. The transducers are silicon capacitive sensors with pressure dependant dielectricum. The pressure sensor contains as dielectricum a small vacuum chamber between the two electrode plates, with the external pressure defining the distance of these plates. Detectors with diaphragms of different pressure sensitivity will be utilized to cover the pressure range to  $\sim 10$  bar. The pressure is derived



as a frequency measurement (within 3–20 kHz range) and the measurements internally compensate for thermal and radiation influences.

#### *5.4. Doppler Wind Experiment*

The primary goal of a Doppler Wind Experiment (DWE) on a Saturn probe would be to measure a vertical profile of the zonal (east-west) winds along the probe descent path. A secondary goal of the DWE is to detect, characterize, and quantify microstructure in the probe descent dynamics, including probe spin, swing, aerodynamic buffeting and atmospheric turbulence, and to detect regions of wind shear and atmospheric wave phenomena. Because of the need for accurate probe and carrier trajectories for making the Doppler wind measurement, the DWE must be closely coordinated with the navigation and radiometric tracking of the carrier, and the probe entry and descent trajectory reconstructions. A Doppler Wind Experiment could be designed to work with a probe DTE communication architecture or a probe-to-relay architecture. Both options include ultra-stable oscillator (USO) requirements and differ only in the angle of entry and DTE geometry requirements. For relay, the system comprises a probe and a carrier USO as part of the probe-carrier communication package. The experiment would benefit from the heritage of both the Galileo and Huygens Doppler Wind Experiments ([Atkinson et al., 1998](#); [Bird et al., 2002](#)).

#### *5.5. Nephelometer*

The composition and precise location of cloud layers in Saturn are largely unknown. They may be composed of ammonia, ammonium hydrosulfide, or simply water. Because of this relative paucity of information on Saturn’s

clouds, the demands we place on a cloud particle sensor, a nephelometer, are significant. Such an instrument would have little heritage in planetary exploration, beyond the one on the Galileo probe. There are similar laser driven, fiber fed nephelometers used in very similar settings on Earth (e.g., [Barkey and Liou, 2001](#); [Barkey et al., 1999](#); [Gayet et al., 1997](#)). However, these were shrouded designs, which is mass-prohibitive on a planetary probe, and they also did not attempt to measure the polarization ratio phase function, because they knew their aerosols were water. The polarization modulation technique that we are proposing was first described by [Hunt and Huffman \(1973\)](#), and has been used in laboratory settings by several groups (e.g., [Kuik et al., 1991](#)). While the precise implementation of the instrument is novel to planetary science, and the polarization modulation technique is also new to an *in situ* instrument, the technology needed to carry out this instrument is all relatively modest. This nephelometer would measure not only the amplitude phase function of the light scattered by the clouds from a laser source on the probe, but also the polarization ratio phase function as well.

#### 5.6. Net Energy Flux Radiometer

A Net Energy Flux Radiometer (NFR) measures the thermal profile and heat budget in the atmosphere. Such a NFR instrument was part of the scientific payload of the Galileo mission ([Sromovsky et al., 1992](#)), which measured the vertical profile of upward and downward radiation fluxes in the region between 0.44 to 14 bar region ([Sromovsky et al., 1998](#)). Radiation was measured in five wavelength bands, 0.3–3.5  $\mu\text{m}$  (total solar radiation), 0.6–3.5  $\mu\text{m}$  (total solar radiation in methane absorption region), 3–500  $\mu\text{m}$  (deposition and loss of thermal radiation), 3.5–5.8  $\mu\text{m}$  (water vapor and cloud struc-

ture), and 14–35  $\mu\text{m}$  (water vapor). On Galileo, NFR found signatures of  $\text{NH}_3$  ice clouds and  $\text{NH}_4\text{SH}$  clouds (Sromovsky et al., 1998), however, water fraction was found to be much lower than solar and no water clouds could be indentified.

## 6. Conclusions

In this paper, we have shown that the *in situ* exploration of Saturn can address two major science themes: the formation history of our solar system and the processes at work in the atmospheres of giant planets. We provided a list of recommended measurements in Saturn’s atmosphere that would allow disentangling between the existing giant planets formation scenarios and the different volatile reservoirs from which the solar system bodies were assembled. Moreover, we illustrated how an entry probe would reveal new insights concerning the vertical structures of temperatures, density, chemical composition and clouds during atmospheric descent. In this context, the top level science goals of a Saturn probe mission would be the determination of:

1. the atmospheric temperature, pressure and mean molecular weight profiles;
2. the abundances of cosmogenically abundant species C, N, S and O;
3. the abundances of chemically inert noble gases He, Ne, Xe, Kr and Ar;
4. the isotopic ratios in hydrogen, oxygen, carbon, nitrogen, He, Ne, Xe, Kr and Ar;
5. the abundances of minor species delivered by vertical mixing (e.g., P, As, Ge) from the deeper troposphere, photochemical species (e.g., hy-

drocarbons, HCN, hydrazine and diphosphine) in the troposphere and exogenic inputs (oxygenated species) in the upper atmosphere;

6. the particle optical properties, size distributions, number and mass densities, opacity, shapes and composition.

Additional *in situ* science measurements aiming at investigating the global electric circuit on Saturn could be also considered (measurement of the Schumann resonances, determination of the vertical profile of conductivity and the spectral power of Saturn lightning at frequencies below the ionospheric cutoff, etc).

We advocated that a Saturn mission incorporating elements of *in situ* exploration should form an essential element of ESA and NASA's future cornerstone missions. We described the concept of a Saturn probe as the next natural step beyond Galileo's *in situ* exploration of Jupiter, and the Cassini spacecraft's orbital reconnaissance of Saturn. Several missions designs have been discussed, all including a spacecraft carrier/orbiter and a probe that would derive from the KRONOS concept previously proposed to ESA ([Marty et al., 2009](#)). International collaborations, in particular between NASA/USA and ESA/Europe may be envisaged in the future to enable the success of a mission devoted to the *in situ* exploration of Saturn.

## Acknowledgements

O.M. acknowledges support from CNES. L.N.F was supported by a Royal Society Research Fellowship at the University of Oxford. P.W. acknowledges support from the Swiss National Science Foundation. O. V. acknowledges support from the KU Leuven IDO project IDO/10/2013 and from the FWO

Postdoctoral Fellowship program. A. S. L and R. H. were supported by the Spanish MICIIN project AYA2012-36666 with FEDER support, PRICIT-S2009/ESP-1496, Grupos Gobierno Vasco IT765-13 and UPV/EHU UFI11/55.

## References

- Alibert, Y., Mousis, O., Benz, W. 2005. On the Volatile Enrichments and Composition of Jupiter. *The Astrophysical Journal* 622, L145-L148.
- Asplund, M., Grevesse, N., Sauval, A. J., Scott, P. 2009. The Chemical Composition of the Sun. *Annual Review of Astronomy and Astrophysics* 47, 481-522.
- Atkinson, D. H., Pollack, J. B., Seiff, A. 1998. The Galileo probe Doppler wind experiment: Measurement of the deep zonal winds on Jupiter. *Journal of Geophysical Research* 103, 22911-22928.
- Atreya, S. K., Mahaffy, P. R., Niemann, H. B., Wong, M. H., Owen, T. C. 2003. Composition and origin of the atmosphere of Jupiter - an update, and implications for the extrasolar giant planets. *Planetary and Space Science* 51, 105-112.
- Atreya, S. K., Wong, M. H., Owen, T. C., Mahaffy, P. R., Niemann, H. B., de Pater, I., Drossart, P., Encrenaz, T. 1999. A comparison of the atmospheres of Jupiter and Saturn: deep atmospheric composition, cloud structure, vertical mixing, and origin. *Planetary and Space Science* 47, 1243-1262.
- Baines, K. H., Carlson, R. W., Kamp, L. W. 2002. Fresh Ammonia Ice

- Clouds in Jupiter. I. Spectroscopic Identification, Spatial Distribution, and Dynamical Implications. *Icarus* 159, 74-94.
- Baldwin, M. P., and 14 colleagues 2001. The quasi-biennial oscillation. *Reviews of Geophysics* 39, 179-229.
- Balsiger, H., and 49 colleagues 2007. Rosina Rosetta Orbiter Spectrometer for Ion and Neutral Analysis. *Space Science Reviews* 128, 745-801.
- Barkey, B., Liou, K. N., Gellerman, W., and Sokolsky, P. 1999. An analog light scattering experiment of hexagonal icelike particles. I. Experimental apparatus and test measurements, *Journal of Atmospheric Science*, Vol. 56(4), 605–612.
- Barkey, B., and Liou, K. N. 2001. Polar nephelometer for light-scattering measurements of ice crystals, *Optics Letters*, 26(4), 2001, 232–234.
- Bézar, B., Lellouch, E., Strobel, D., Maillard, J.-P., Drossart, P. 2002. Carbon Monoxide on Jupiter: Evidence for Both Internal and External Sources. *Icarus* 159, 95-111.
- Bird, M. K., and 12 colleagues 2002. The Huygens Doppler Wind Experiment - Titan Winds Derived from Probe Radio Frequency Measurements. *Space Science Reviews* 104, 613-640.
- Boss, A. P. 2001. Formation of Planetary-Mass Objects by Protostellar Collapse and Fragmentation. *The Astrophysical Journal* 551, L167-L170.
- Boss, A. P. 1997. Giant planet formation by gravitational instability. *Science* 276, 1836-1839.

- Briggs, F. H., Sackett, P. D. 1989. Radio observations of Saturn as a probe of its atmosphere and cloud structure. *Icarus* 80, 77-103.
- Briois, C., and 22 colleagues 2013. Dust OrbiTrap Sensor (DOTS) for In-Situ Analysis of Airless Planetary Bodies. Lunar and Planetary Institute Science Conference Abstracts 44, 2888.
- Cavalié, T., and 10 colleagues 2014. The first submillimeter observation of CO in the stratosphere of Uranus. *Astronomy and Astrophysics*, in press.
- Cavalié, T., Hartogh, P., Billebaud, F., Dobrijevic, M., Fouchet, T., Lellouch, E., Encrenaz, T., Brillet, J., Moriarty-Schieven, G. H. 2010. A cometary origin for CO in the stratosphere of Saturn?. *Astronomy and Astrophysics* 510, A88.
- Cavalié, T., Billebaud, F., Dobrijevic, M., Fouchet, T., Lellouch, E., Encrenaz, T., Brillet, J., Moriarty-Schieven, G. H., Wouterloot, J. G. A., Hartogh, P. 2009. First observation of CO at 345 GHz in the atmosphere of Saturn with the JCMT: New constraints on its origin. *Icarus* 203, 531-540.
- Chambers, J. E., Wetherill, G. W. 2001. Planets in the asteroid belt. *Meteoritics and Planetary Science* 36, 381-399.
- Chick, K. M., Cassen, P. 1997. Thermal Processing of Interstellar Dust Grains in the Primitive Solar Environment. *The Astrophysical Journal* 477, 398.
- Conrath, B. J., Gautier, D. 2000. Saturn Helium Abundance: A Reanalysis of Voyager Measurements. *Icarus* 144, 124-134.

- de Graauw, T., and 18 colleagues 1997. First results of ISO-SWS observations of Saturn: detection of CO<sub>2</sub>, CH<sub>3</sub>C<sub>2</sub>H, C<sub>4</sub>H<sub>2</sub> and tropospheric H<sub>2</sub>O. *Astronomy and Astrophysics* 321, L13-L16.
- Del Genio, A. D., Achterberg, R. K., Baines, K. H., Flasar, F. M., Read, P. L., Sánchez-Lavega, A., Showman, A. P. 2009. Saturn Atmospheric Structure and Dynamics. *Saturn from Cassini-Huygens* 113.
- Durry, G., and 12 colleagues 2010. Near infrared diode laser spectroscopy of C<sub>2</sub>H<sub>2</sub>, H<sub>2</sub>O, CO<sub>2</sub> and their isotopologues and the application to TDLAS, a tunable diode laser spectrometer for the martian PHOBOS-GRUNT space mission. *Applied Physics B: Lasers and Optics* 99, 339-351.
- Durry, G., Hauchecorne, A., Ovarlez, J., Ovarlez, H., Pouchet, I., Zeninari, V., Parvitte, B. 2002. In situ Measurement of H<sub>2</sub>O and CH<sub>4</sub> with Telecommunication Laser Diodes in the Lower Stratosphere: Dehydration and Indication of a Tropical Air Intrusion at Mid-Latitudes. *Journal of Atmospheric Chemistry* 43(3), 175-194.
- Ellerby, D., Venkatapathy, E., Stackpoole, M., Chinnapongse, R., Beerman, A., Feldman, J., Peterson, K., Prahbu, D., and Munk, M. 2013. Woven Thermal Protection System (WTPS) a Novel Approach to Meet NASA's Most Demanding Reentry Missions. 10th International Planetary Probe Workshop, session 8B: Cross Cutting Technologies III.
- Durry, G., Hauchecorne, A., Ovarlez, J., Ovarlez, H., Pouchet, I., Zeninari, V., Parvitte, B. 2002. In situ Measurement of H<sub>2</sub>O and CH<sub>4</sub> with Telecommunication Laser Diodes in the Lower Stratosphere: Dehydration



- and Indication of a Tropical Air Intrusion at Mid-Latitudes. *Journal of Atmospheric Chemistry* 43(3), 175-194.
- Fegley, B., Jr., Lodders, K. 1994. Chemical models of the deep atmospheres of Jupiter and Saturn. *Icarus* 110, 117-154.
- Fletcher, L. N., and 17 colleagues 2012. Sub-millimetre spectroscopy of Saturn's trace gases from Herschel/SPIRE. *Astronomy and Astrophysics* 539, A44.
- Fletcher, L. N., Baines, K. H., Momary, T. W., Showman, A. P., Irwin, P. G. J., Orton, G. S., Roos-Serote, M., Merlet, C. 2011. Saturn's tropospheric composition and clouds from Cassini/VIMS 4.6-5.1  $\mu\text{m}$  nightside spectroscopy. *Icarus* 214, 510-533.
- Fletcher, L. N., Orton, G. S., Teanby, N. A., Irwin, P. G. J. 2009a. Phosphine on Jupiter and Saturn from Cassini/CIRS. *Icarus* 202, 543-564.
- Fletcher, L. N., Orton, G. S., Teanby, N. A., Irwin, P. G. J., Bjoraker, G. L. 2009b. Methane and its isotopologues on Saturn from Cassini/CIRS observations. *Icarus* 199, 351-367.
- Feuchtgruber, H., Lellouch, E., de Graauw, T., Bézard, B., Encrenaz, T., Griffin, M. 1997. External supply of oxygen to the atmospheres of the giant planets. *Nature* 389, 159-162.
- Folkner, W. M., Woo, R., Nandi, S. 1998. Ammonia abundance in Jupiter's atmosphere derived from the attenuation of the Galileo probe's radio signal. *Journal of Geophysical Research* 103, 22847-22856.

- Fortney, J. J., Nettelmann, N. 2010. The Interior Structure, Composition, and Evolution of Giant Planets. *Space Science Reviews* 152, 423-447.
- Fortney, J. J., Hubbard, W. B. 2003. Phase separation in giant planets: inhomogeneous evolution of Saturn. *Icarus* 164, 228-243.
- Fouchet, T., Moses, J. I., Conrath, B. J. 2009. Saturn: Composition and Chemistry. *Saturn from Cassini-Huygens* 83.
- Fouchet, T., Guerlet, S., Strobel, D. F., Simon-Miller, A. A., Bézard, B., Flasar, F. M. 2008. An equatorial oscillation in Saturn's middle atmosphere. *Nature* 453, 200-202.
- Fouchet, T., Irwin, P. G. J., Parrish, P., Calcutt, S. B., Taylor, F. W., Nixon, C. A., Owen, T. 2004. Search for spatial variation in the jovian  $^{15}\text{N}/^{14}\text{N}$  ratio from Cassini/CIRS observations. *Icarus* 172, 50-58.
- Fouchet, T., Lellouch, E., Bézard, B., Encrenaz, T., Drossart, P., Feuchtgruber, H., de Graauw, T. 2000. ISO-SWS Observations of Jupiter: Measurement of the Ammonia Tropospheric Profile and of the  $^{15}\text{N}/^{14}\text{N}$  Isotopic Ratio. *Icarus* 143, 223-243.
- Friedson, A. J. 1999. New Observations and Modelling of a QBO-Like Oscillation in Jupiter's Stratosphere. *Icarus* 137, 34-55.
- Fulchignoni, M., et al., 2002, The Characterisation of Titan's Atmospheric Physical Properties by the Huygens Atmospheric Structure Instrument (HASI), *Space Science Reviews*, 104(1), 397-434

- García-Melendo, E., Pérez-Hoyos, S., Sánchez-Lavega, A., Hueso, R. 2011. Saturn's zonal wind profile in 2004-2009 from Cassini ISS images and its long-term variability. *Icarus* 215, 62-74.
- Gautier, D., Hersant, F., Mousis, O., Lunine, J. I. 2001. Enrichments in Volatiles in Jupiter: A New Interpretation of the Galileo Measurements. *The Astrophysical Journal* 550, L227-L230.
- Gayet, J. F., Crépel, O., Fournol, J. F., and Oshchepkov, S. 1997. A new airborne polar Nephelometer for the measurements of optical and microphysical cloud properties. Part I: Theoretical design, *Annales Geophysicae*, 15(4), 451–459.
- Goesmann, F., Rosenbauer, H., Roll, R., Szopa, C., Raulin, F., Sternberg, R., Israel, G., Meierhenrich, U., Thiemann, W., Munoz-Caro, G. 2007. Cosac, The Cometary Sampling and Composition Experiment on Philae. *Space Science Reviews* 128, 257-280.
- Gomes, R., Levison, H. F., Tsiganis, K., Morbidelli, A. 2005. Origin of the cataclysmic Late Heavy Bombardment period of the terrestrial planets. *Nature* 435, 466-469.
- Guerlet, S., Fouchet, T., Bézard, B., Flasar, F. M., Simon-Miller, A. A. 2011. Evolution of the equatorial oscillation in Saturn's stratosphere between 2005 and 2010 from Cassini/CIRS limb data analysis. *Geophysical Research Letters* 38, s9201.
- Guillot, T., Hueso, R. 2006. The composition of Jupiter: sign of a (relatively)

- late formation in a chemically evolved protosolar disc. *Monthly Notices of the Royal Astronomical Society* 367, L47-L51.
- Guillot, T. 2004. Probing the Giant Planets. *Physics Today* 57, 040000-70.
- Guillot, T. 1999. A comparison of the interiors of Jupiter and Saturn. *Planetary and Space Science* 47, 1183-1200.
- Hartogh, P., and 11 colleagues 2011. Direct detection of the Enceladus water torus with Herschel. *Astronomy and Astrophysics* 532, L2.
- Helled, R., Guillot, T. 2013. Interior Models of Saturn: Including the Uncertainties in Shape and Rotation. *The Astrophysical Journal* 767, 113.
- Hersant, F., Gautier, D., Tobie, G., Lunine, J. I. 2008. Interpretation of the carbon abundance in Saturn measured by Cassini. *Planetary and Space Science* 56, 1103-1111.
- Hersant, F., Gautier, D., Lunine, J. I. 2004. Enrichment in volatiles in the giant planets of the Solar System. *Planetary and Space Science* 52, 623-641.
- Hickey, M. P., Walterscheid, R. L., Schubert, G. 2000. Gravity Wave Heating and Cooling in Jupiter's Thermosphere. *Icarus* 148, 266-281.
- Hörst, S. M., and 12 colleagues 2012. Formation of Amino Acids and Nucleotide Bases in a Titan Atmosphere Simulation Experiment. *Astrobiology* 12, 809-817.
- Hu, Q., Noll, R. J., Li, H., Makarov, A., Hardman, M., and Cooks, R. G. 2005. The Orbitrap: a new mass spectrometer. *Journal of Mass Spectrometry*, 40(4), 430-443.

- Hunt, A. J., and Huffman, D. R. 1973. The Clouds of Venus-An Experimental Multiple Scattering Study of the Polarization, Bulletin of the American Astronomical Society, Vol. 5, p.416.
- Ida, S., Lin, D. N. C. 2004. Toward a Deterministic Model of Planetary Formation. I. A Desert in the Mass and Semimajor Axis Distributions of Extrasolar Planets. The Astrophysical Journal 604, 388-413.
- Janhunen, P., Lebreton, J.-P., Merikallio, S., Paton, M., Mengali, G., and Quarta, A. 2014. Fast E-sail Uranus entry probe mission. Planetary and Space Science, submitted.
- Johnson, T. V., Mousis, O., Lunine, J. I., Madhusudhan, N. 2012. Planetesimal Compositions in Exoplanet Systems. The Astrophysical Journal 757, 192.
- Kouchi, A., Yamamoto, T., Kozasa, T., Kuroda, T., Greenberg, J. M. 1994. Conditions for condensation and preservation of amorphous ice and crystallinity of astrophysical ices. Astronomy and Astrophysics 290, 1009-1018.
- Kuik, F., Stammes, P., Streekstra, M. L., and Hovenier, J. W. 1991. Measurements of scattering matrices of water droplets and ice crystals, in *Digest of the ICO topical meeting on atmospheric, volume and surface scattering, and propagation*, Florence 1991, pages 397–401.
- Laraia, A. L., Ingersoll, A. P., Janssen, M. A., Gulkis, S., Oyafuso, F., Allison, M. 2013. Analysis of Saturn's thermal emission at 2.2-cm wavelength: Spatial distribution of ammonia vapor. Icarus 226, 641-654.

- Le Barbu, T., Vinogradov, I., Durry, G., Korablev, O., Chassefière, E., Bertaux, J.-L. 2004. Tdlas, a diode laser sensor for the in situ monitoring of H<sub>2</sub>O and CO<sub>2</sub> isotopes. 35th COSPAR Scientific Assembly 35, 2115.
- Leconte, J., Chabrier, G. 2013. Layered convection as the origin of Saturn's luminosity anomaly. *Nature Geoscience* 6, 347-350.
- Leconte, J., Chabrier, G. 2012. A new vision of giant planet interiors: Impact of double diffusive convection. *Astronomy and Astrophysics* 540, A20.
- Lellouch, E., Bézard, B., Fouchet, T., Feuchtgruber, H., Encrenaz, T., de Graauw, T. 2001. The deuterium abundance in Jupiter and Saturn from ISO-SWS observations. *Astronomy and Astrophysics* 370, 610-622.
- Lin, D. N. C., Papaloizou, J. 1986. On the tidal interaction between protoplanets and the protoplanetary disk. III - Orbital migration of protoplanets. *The Astrophysical Journal* 309, 846-857.
- Lodders, K., Palme, H., Gail, H.-P. 2009. Abundances of the Elements in the Solar System. *Landolt Börnstein* 44.
- Lodders, K. 2004. Jupiter Formed with More Tar than Ice. *The Astrophysical Journal* 611, 587-597.
- Magalhães, J. A., Seiff, A., Young, R. E. 2002. The Stratification of Jupiter's Troposphere at the Galileo Probe Entry Site. *Icarus* 158, 410-433.
- Mahaffy, P. R., and 84 colleagues 2012. The Sample Analysis at Mars Investigation and Instrument Suite. *Space Science Reviews* 170, 401-478.

- Mahaffy, P. R., Niemann, H. B., Alpert, A., Atreya, S. K., Demick, J., Donahue, T. M., Harpold, D. N., Owen, T. C. 2000. Noble gas abundance and isotope ratios in the atmosphere of Jupiter from the Galileo Probe Mass Spectrometer. *Journal of Geophysical Research* 105, 15061-15072.
- Makarov, A., 2000. Electrostatic Axially Harmonic Orbital Trapping: A High-Performance Technique of Mass Analysis. *Anal. Chem.* 72, 1156-1162.
- Mandt, K. E., Waite, J. H., Lewis, W., Magee, B., Bell, J., Lunine, J., Mousis, O., Cordier, D. 2009. Isotopic evolution of the major constituents of Titan's atmosphere based on Cassini data. *Planetary and Space Science* 57, 1917-1930.
- Marley, M. S., Ackerman, A. S., Cuzzi, J. N., Kitzmann, D. 2013. Clouds and Hazes in Exoplanet Atmospheres. ArXiv e-prints arXiv:1301.5627.
- Marty, B., Chaussidon, M., Wiens, R. C., Jurewicz, A. J. G., Burnett, D. S. 2011. A  $^{15}\text{N}$ -Poor Isotopic Composition for the Solar System As Shown by Genesis Solar Wind Samples. *Science* 332, 1533.
- Marty, B., and 56 colleagues 2009. Kronos: exploring the depths of Saturn with probes and remote sensing through an international mission. *Experimental Astronomy* 23, 947-976.
- McKeegan, Kevin D., and 46 colleagues 2006. Isotopic Compositions of Cometary Matter Returned by Stardust. *Science* 314, 1724.
- Miller, T. A., Wooldridge, M. S., Bozzelli, J. W. 2004. Computational modeling of the  $\text{SiH}_3 + \text{O}_2$  reaction and silane combustion. *Combustion and Flame* 137, 73-92.

- Mizuno, H. 1980. Formation of the Giant Planets. *Progress of Theoretical Physics* 64, 544-557.
- Mordasini, C., Alibert, Y., Klahr, H., Henning, T. 2012. Characterization of exoplanets from their formation. I. Models of combined planet formation and evolution. *Astronomy and Astrophysics* 547, A111.
- Moses, J. I., Fouchet, T., Bézard, B., Gladstone, G. R., Lellouch, E., Feuchgruber, H. 2005. Photochemistry and diffusion in Jupiter's stratosphere: Constraints from ISO observations and comparisons with other giant planets. *Journal of Geophysical Research (Planets)* 110, 8001.
- Mousis, O., Lunine, J. I., Madhusudhan, N., Johnson, T. V. 2012. Nebular Water Depletion as the Cause of Jupiter's Low Oxygen Abundance. *The Astrophysical Journal* 751, L7.
- Mousis, O., Marboeuf, U., Lunine, J. I., Alibert, Y., Fletcher, L. N., Orton, G. S., Pauzat, F., Ellinger, Y. 2009. Determination of the Minimum Masses of Heavy Elements in the Envelopes of Jupiter and Saturn. *The Astrophysical Journal* 696, 1348-1354.
- Mousis, O., Alibert, Y., Benz, W. 2006. Saturn's internal structure and carbon enrichment. *Astronomy and Astrophysics* 449, 411-415.
- Nagy, A. F., Kliore, A. J., Mendillo, M., Miller, S., Moore, L., Moses, J. I., Müller-Wodarg, I., Shemansky, D. 2009. Upper Atmosphere and Ionosphere of Saturn. *Saturn from Cassini-Huygens* 181.
- Nettelmann, N., Püstow, R., Redmer, R. 2013. Saturn layered structure and homogeneous evolution models with different EOSs. *Icarus* 225, 548-557.



- Niemann, H. B., Atreya, S. K., Demick, J. E., Gautier, D., Haberman, J. A., Harpold, D. N., Kasprzak, W. T., Lunine, J. I., Owen, T. C., Raulin, F. 2010. Composition of Titan's lower atmosphere and simple surface volatiles as measured by the Cassini-Huygens probe gas chromatograph mass spectrometer experiment. *Journal of Geophysical Research (Planets)* 115, 12006.
- Niemann, H. B., and 17 colleagues 2005. The abundances of constituents of Titan's atmosphere from the GCMS instrument on the Huygens probe. *Nature* 438, 779-784.
- Niemann, H. B., and 18 colleagues 2002. The Gas Chromatograph Mass Spectrometer for the Huygens Probe. *Space Science Reviews* 104, 551-590.
- Niemann, H. B., and 11 colleagues 1998. The composition of the Jovian atmosphere as determined by the Galileo probe mass spectrometer. *Journal of Geophysical Research* 103, 22831-22846.
- Niemann, H. B., and 12 colleagues 1996. The Galileo Probe Mass Spectrometer: Composition of Jupiter's Atmosphere. *Science* 272, 846-849.
- Niemann, H. B., Harpold, D. N., Atreya, S. K., Carignan, G. R., Hunten, D. M., Owen, T. C. 1992. Galileo Probe Mass Spectrometer experiment. *Space Science Reviews* 60, 111-142.
- Noll, K. S., Geballe, T. R., Knacke, R. F. 1995. Detection of H<sub>2</sub> 18O in Jupiter. *The Astrophysical Journal* 453, L49.
- Notesco, G., Bar-Nun, A. 2005. A ~25 K temperature of formation for the submicron ice grains which formed comets. *Icarus* 175, 546-550.

- Okumura, D., Toyoda, M., Ishihara, M., Katakuse, I. 2004. A compact sector-type multi-turn time-of-flight mass spectrometer 'MULTUM II'. *Nuclear Instruments and Methods in Physics Research A* 519, 331-337.
- Orton, G. S., and 27 colleagues 2008. Semi-annual oscillations in Saturn's low-latitude stratospheric temperatures. *Nature* 453, 196-199.
- Orton, G. S., and 16 colleagues 1998. Characteristics of the Galileo probe entry site from Earth-based remote sensing observations. *Journal of Geophysical Research* 103, 22791-22814.
- Owen, T., Mahaffy, P. R., Niemann, H. B., Atreya, S., Wong, M. 2001. Protosolar Nitrogen. *The Astrophysical Journal* 553, L77-L79.
- Owen, T., Mahaffy, P., Niemann, H. B., Atreya, S., Donahue, T., Bar-Nun, A., de Pater, I. 1999. A low-temperature origin for the planetesimals that formed Jupiter. *Nature* 402, 269-270.
- Pasek, M. A., Milsom, J. A., Ciesla, F. J., Lauretta, D. S., Sharp, C. M., Lunine, J. I. 2005. Sulfur chemistry with time-varying oxygen abundance during Solar System formation. *Icarus* 175, 1-14.
- Pollack, J. B., Hubickyj, O., Bodenheimer, P., Lissauer, J. J., Podolak, M., Greenzweig, Y. 1996. Formation of the Giant Planets by Concurrent Accretion of Solids and Gas. *Icarus* 124, 62-85.
- Ragent, B., Colburn, D. S., Rages, K. A., Knight, T. C. D., Avrin, P., Orton, G. S., Yanamandra-Fisher, P. A., Grams, G. W. 1998. The clouds of Jupiter: Results of the Galileo Jupiter mission probe nephelometer experiment. *Journal of Geophysical Research* 103, 22891-22910.

- Read, P. L., Conrath, B. J., Fletcher, L. N., Gierasch, P. J., Simon-Miller, A. A., Zuchowski, L. C. 2009. Mapping potential vorticity dynamics on saturn: Zonal mean circulation from Cassini and Voyager data. *Planetary and Space Science* 57, 1682-1698.
- Reuter, D. C., and 10 colleagues 2007. Jupiter Cloud Composition, Stratification, Convection, and Wave Motion: A View from New Horizons. *Science* 318, 223.
- Roos-Serote, M., Vasavada, A. R., Kamp, L., Drossart, P., Irwin, P., Nixon, C., Carlson, R. W. 2000. Proximate humid and dry regions in Jupiter's atmosphere indicate complex local meteorology. *Nature* 405, 158-160.
- Rousselot, P., and 11 colleagues 2014. Toward a Unique Nitrogen Isotopic Ratio in Cometary Ices. *The Astrophysical Journal* 780, L17.
- Sánchez-Lavega, A., and 24 colleagues 2008. Depth of a strong jovian jet from a planetary-scale disturbance driven by storms. *Nature* 451, 437-440.
- Sanchez-Lavega, A., Colas, F., Lecacheux, J., Laques, P., Parker, D., Miyazaki, I. 1991. The Great White SPOT and disturbances in Saturn's equatorial atmosphere during 1990. *Nature* 353, 397-401.
- Saumon, D., Guillot, T. 2004. Shock Compression of Deuterium and the Interiors of Jupiter and Saturn. *The Astrophysical Journal* 609, 1170-1180.
- Scherer, S., Altwegg, K., Balsiger, H., Fischer, J., Jäckel, A., Korth, A., Mildner, M., Piazza, D., Reme, H., Wurz, P. 2006. A novel principle for an ion mirror design in time-of-flight mass spectrometry. *International Journal of Mass Spectrometry* 251, 73-81.

- Seiff, A., Kirk, D. B., Knight, T. C. D., Young, R. E., Mihalov, J. D., Young, L. A., Milos, F. S., Schubert, G., Blanchard, R. C., Atkinson, D. 1998. Thermal structure of Jupiter's atmosphere near the edge of a 5- $\mu$ m hot spot in the north equatorial belt. *Journal of Geophysical Research* 103, 22857-22890.
- Showman, A. P., Dowling, T. E. 2000. Nonlinear Simulations of Jupiter's 5-Micron Hot Spots. *Science* 289, 1737-1740.
- Simon-Miller, A. A., Conrath, B., Gierasch, P. J., Beebe, R. F. 2000. A Detection of Water Ice on Jupiter with Voyager IRIS. *Icarus* 145, 454-461.
- Stevenson, D. J., Lunine, J. I. 1988. Rapid formation of Jupiter by diffuse redistribution of water vapor in the solar nebula. *Icarus* 75, 146-155.
- Stevenson, D. J., Salpeter, E. E. 1977a. The dynamics and helium distribution in hydrogen-helium fluid planets. *The Astrophysical Journal Supplement Series* 35, 239-261.
- Stevenson, D. J., Salpeter, E. E. 1977b. The phase diagram and transport properties for hydrogen-helium fluid planets. *The Astrophysical Journal Supplement Series* 35, 221-237.
- Sromovsky, L. A., Baines, K. H., Fry, P. M. 2013. Saturn's Great Storm of 2010-2011: Evidence for ammonia and water ices from analysis of VIMS spectra. *Icarus* 226, 402-418.
- Sromovsky, L. A., Fry, P. M. 2010. The source of widespread 3- $\mu$ m absorption in Jupiter's clouds: Constraints from 2000 Cassini VIMS observations. *Icarus* 210, 230-257.

- Sromovsky, L. A., Collard, A. D., Fry, P. M., Orton, G. S., Lemmon, M. T., Tomasko, M. G., Freedman, R. S. 1998. Galileo probe measurements of thermal and solar radiation fluxes in the Jovian atmosphere. *Journal of Geophysical Research* 103, 2929.
- Sromovsky, L. A., Best, F. A., Revercomb, H. E., and Hayden, J., 1992, Galileo Net Flux Radiometer experiment, *Space Science Reviews* 60(1-4), 1992, 233–262.
- Teanby, N. A., Fletcher, L. N., Irwin, P. G. J., Fouchet, T., Orton, G. S. 2006. New upper limits for hydrogen halides on Saturn derived from Cassini-CIRS data. *Icarus* 185, 466-475.
- Tomasko, M. G., and 13 colleagues 2002. The Descent Imager/Spectral Radiometer (DISR) Experiment on the Huygens Entry Probe of Titan. *Space Science Reviews* 104, 469-551.
- Toyoda, M., Okumura, D., Ishihara, M., Katakuse I. 2003. Multi-turn time-of-flight mass spectrometers with electrostatic sectors. *J. Mass Spectrom.* 38, 1125-1142.
- Twarowski, A. 1995. Reduction of a phosphorus oxide and acid reaction set. *Combustion and Flame* 102, 41-54.
- Venot, O., Hébrard, E., Agúndez, M., Dobrijevic, M., Selsis, F., Hersant, F., Iro, N., Bounaceur, R. 2012. A chemical model for the atmosphere of hot Jupiters. *Astronomy and Astrophysics* 546, A43.
- Visscher, C., Moses, J. I., Saslow, S. A. 2010. The deep water abundance

- on Jupiter: New constraints from thermochemical kinetics and diffusion modeling. *Icarus* 209, 602-615.
- Visscher, C., Fegley, B., Jr. 2005. Chemical Constraints on the Water and Total Oxygen Abundances in the Deep Atmosphere of Saturn. *The Astrophysical Journal* 623, 1221-1227.
- von Zahn, U., Hunten, D. M., Lehmacher, G. 1998. Helium in Jupiter's atmosphere: Results from the Galileo probe helium interferometer experiment. *Journal of Geophysical Research* 103, 22815-22830.
- von Zahn, U., and Hunten, D. M., 1992, The Jupiter Helium Interferometer experiment on the Galileo entry probe, *Space Science Reviews*, 60(1-4), 263-281.
- Waite, H. Jr., and 10 colleagues 2012. A Neutral Gas Investigation of Origins (ANGIO), submitted to NASA AO NNH12ZDA006O-JUICE, Jupiter Icy Moons Explorer Instrument.
- Ward, W. R. 1997. Protoplanet Migration by Nebula Tides. *Icarus* 126, 261-281.
- Watkins, C., Cho, J. Y.-K. 2013. The vertical structure of Jupiter's equatorial zonal wind above the cloud deck, derived using mesoscale gravity waves. *Geophysical Research Letters* 40, 472-476.
- Webster, C. R., and 13 colleagues 2013. Isotope Ratios of H, C, and O in CO<sub>2</sub> and H<sub>2</sub>O of the Martian Atmosphere. *Science* 341, 260-263.

- Webster, C. R., Mahaffy, P. R. 2011. Determining the local abundance of Martian methane and its  $^{13}\text{C}/^{12}\text{C}$  and D/H isotopic ratios for comparison with related gas and soil analysis on the 2011 Mars Science Laboratory (MSL) mission. *Planetary and Space Science* 59, 271-283.
- Weidenschilling, S. J., Lewis, J. S. 1973. Atmospheric and cloud structures of the Jovian planets. *Icarus* 20, 465-476.
- West, R. A., Baines, K. H., Karkoschka, E., Sánchez-Lavega, A. 2009. Clouds and Aerosols in Saturn's Atmosphere. *Saturn from Cassini-Huygens* 161.
- West, R. A., Baines, K. H., Friedson, A. J., Banfield, D., Ragent, B., Taylor, F. W. 2004. Jovian clouds and haze. *Jupiter. The Planet, Satellites and Magnetosphere* 79-104.
- Wilson, H. F., Militzer, B. 2012. Rocky Core Solubility in Jupiter and Giant Exoplanets. *Physical Review Letters* 108, 111101.
- Wilson, H., Militzer, B. 2011. Solubility and erosion of icy cores in giant planets. *APS Meeting Abstracts* 31010.
- Wilson, H. F., Militzer, B. 2010. Sequestration of Noble Gases in Giant Planet Interiors. *Physical Review Letters* 104, 121101.
- Wong, M. H. 2009. Comment on "Transport of nonmethane hydrocarbons to Jupiter's troposphere by descent of smog particles" by Donald M. Hunten [*Icarus* 194 (2008) 616-622]. *Icarus* 199, 231-235.
- Wong, M. H., Mahaffy, P. R., Atreya, S. K., Niemann, H. B., Owen, T. C.

2004. Updated Galileo probe mass spectrometer measurements of carbon, oxygen, nitrogen, and sulfur on Jupiter. *Icarus* 171, 153-170.
- Wright, I. P., and 19 colleagues 2007. Ptolemy an Instrument to Measure Stable Isotopic Ratios of Key Volatiles on a Cometary Nucleus. *Space Science Reviews* 128, 363-381.
- Yelle, R. V., Miller, S. 2004. Jupiter's thermosphere and ionosphere. *Jupiter. The Planet, Satellites and Magnetosphere* 185-218.
- Zarnecki, J. C., Ferri, F., Hathi, B., Leese, M. R., Ball, A. J., Colombatti, G., and Fulchignoni, M., 2004, In-flight performance of the HASI servo accelerometer and implications for results at Titan, in *Proceedings of the International Workshop Planetary Probe Atmospheric Entry and Descent Trajectory Analysis and Science*, ESA SP-544, 2004, 71-76.



Table 1: Compositions of the atmospheres of Jupiter and Saturn (major volatiles)

Species	Jupiter			Saturn		
	X/H <sub>2</sub>	$\Delta(X/H_2)$	Reference	X/H <sub>2</sub>	$\Delta(X/H_2)$	Reference
CH <sub>4</sub>	$2.37 \times 10^{-03}$	$5.70 \times 10^{-04}$	Wong et al. (2004)	$5.33 \times 10^{-3}$	$0.23 \times 10^{-3}$	Fletcher et al. (2009b)
NH <sub>3</sub>	$6.64 \times 10^{-04}$	$2.54 \times 10^{-04}$	Wong et al. (2004)	$1.04\text{--}5.78 \times 10^{-4}$	–	Fletcher et al. (2011)
H <sub>2</sub> O <sup>(a)</sup>	$4.90 \times 10^{-04}$	$1.60 \times 10^{-04}$	Wong et al. (2004)	$2.0 \times 10^{-07}$	–	de Graauw et al. (1997)
PH <sub>3</sub>	$2.11 \times 10^{-06}$	$1.00 \times 10^{-07}$	Fletcher et al. (2009a)	$7.30 \times 10^{-06}$	$0.48 \times 10^{-06}$	Fletcher et al. (2009a)
H <sub>2</sub> S	$8.90 \times 10^{-05}$	$2.10 \times 10^{-05}$	Wong et al. (2004)	$3.90 \times 10^{-04}$	–	Briggs and Sackett (1989)
He	$1.36 \times 10^{-01}$	$2.70 \times 10^{-03}$	von Zahn et al. (1998)	$1.35 \times 10^{-1}$	$0.25 \times 10^{-1}$	Conrath and Gautier (2000)
Ne <sup>(b)</sup>	$3.0 \times 10^{-05}$	–	Niemann et al. (1998)	–	–	–
Ar	$1.82 \times 10^{-05}$	$3.60 \times 10^{-06}$	Mahaffy et al. (2000)	–	–	–
Kr	$9.30 \times 10^{-09}$	$1.70 \times 10^{-09}$	Mahaffy et al. (2000)	–	–	–
Xe	$8.90 \times 10^{-10}$	$1.70 \times 10^{-10}$	Mahaffy et al. (2000)	–	–	–

<sup>(a)</sup>This is a lower limit; <sup>(b)</sup>this is an upper limit.

Table 2: Isotopic ratios in Jupiter and Saturn

Isotopic ratio	Jupiter			Saturn		
	$\eta$	$\Delta\eta$	Reference	$\eta$	$\Delta\eta$	Reference
D/H (in H <sub>2</sub> )	$2.60 \times 10^{-5}$	$0.70 \times 10^{-5}$	Niemann et al. (1998)	$1.70 \times 10^{-5}$	${}^{+0.75}_{-0.45} \times 10^{-05}$	Lellouch et al. (2001)
<sup>3</sup> He/ <sup>4</sup> He	$1.66 \times 10^{-4}$	$0.05 \times 10^{-4}$	Niemann et al. (1998)	—	—	—
<sup>12</sup> C/ <sup>13</sup> C (in CH <sub>4</sub> )	92.6	${}^{+4.5}_{-4.1}$	Niemann et al. (1996)	91.8	${}^{+8.4}_{-7.8}$	Fletcher et al. (2009b)
<sup>14</sup> N/ <sup>15</sup> N (in NH <sub>3</sub> )	434.8	${}^{+65}_{-50}$	Wong et al. (2004)	—	—	—
<sup>20</sup> Ne/ <sup>22</sup> Ne	13.0	2.0	Mahaffy et al. (2000)	—	—	—
<sup>36</sup> Ar/ <sup>38</sup> Ar	5.6	0.25	Mahaffy et al. (2000)	—	—	—
<sup>128</sup> Xe/total Xe	0.018	0.002	Atreya et al. (2003)	—	—	—
<sup>129</sup> Xe/total Xe	0.285	0.021	Atreya et al. (2003)	—	—	—
<sup>130</sup> Xe/total Xe	0.038	0.005	Atreya et al. (2003)	—	—	—
<sup>131</sup> Xe/total Xe	0.203	0.018	Atreya et al. (2003)	—	—	—
<sup>132</sup> Xe/total Xe	0.290	0.020	Atreya et al. (2003)	—	—	—
<sup>134</sup> Xe/total Xe	0.091	0.007	Atreya et al. (2003)	—	—	—
<sup>136</sup> Xe/total Xe	0.076	0.009	Atreya et al. (2003)	—	—	—

Table 3: Enrichments in Jupiter and Saturn relatives to Protosun

Species	Jupiter		Saturn	
	E	$\Delta E^{(a)}$	E	$\Delta E^{(a)}$
C	4.40	1.14	9.90	1.05
N	4.18	2.08	0.53–4.07	–
O <sup>(b)</sup>	0.42	0.15	$\sim 10^{-4}$	–
P	3.34	0.36	11.54	1.35
S	2.94	0.70	15.87	–
He	0.72	0.04	0.71	0.14
Ne <sup>(c)</sup>	0.12	–	–	–
Ar	2.62	0.86	–	–
Kr	2.23	0.61	–	–
Xe	2.18	0.61	–	–

<sup>(a)</sup>Error is defined as  $(\Delta E/E)^2 = (\Delta X/X_{\text{planet}})^2 + (\Delta X/X_{\text{Protosun}})^2$ ; <sup>(b)</sup>this is a lower limit; <sup>(c)</sup>this is an upper limit.

Table 4: Elemental abundances in the Sun and Protosun

Element	Solar dex	Protosolar dex	$\Delta$ dex	Protosolar X/H <sub>2</sub>	$\Delta(X/H_2)$
C	8.39	8.43	0.04	$5.38 \times 10^{-04}$	$5.19 \times 10^{-05}$
N	7.86	7.90	0.12	$1.59 \times 10^{-04}$	$5.06 \times 10^{-05}$
O	8.73	8.77	0.07	$1.18 \times 10^{-03}$	$2.06 \times 10^{-04}$
P	5.46	5.50	0.04	$6.32 \times 10^{-07}$	$6.10 \times 10^{-08}$
S	7.14	7.18	0.01	$3.03 \times 10^{-05}$	$7.05 \times 10^{-07}$
He	10.93	10.98	0.02	$1.89 \times 10^{-01}$	$8.90 \times 10^{-03}$
Ne	8.05	8.09	0.10	$2.46 \times 10^{-04}$	$6.37 \times 10^{-05}$
Ar	6.50	6.54	0.10	$6.93 \times 10^{-06}$	$1.80 \times 10^{-06}$
Kr	3.28	3.32	0.08	$4.18 \times 10^{-09}$	$8.45 \times 10^{-10}$
Xe	2.27	2.31	0.08	$4.08 \times 10^{-10}$	$8.26 \times 10^{-11}$

Data from [Lodders et al. \(2009\)](#) with values of corrections for protosolar abundances (+0.05 dex (He) and +0.04 dex (others)) taken from [Asplund et al. \(2009\)](#).

Table 5: Measurement requirements

Instrument	Measurement
Mass spectrometer	Elemental and chemical composition Isotopic composition High molecular mass organics
Helium abundance detector	Accurate He/H <sub>2</sub> ratio
Atmospheric Structure Instrument	Pressure, temperature, density, molecular weight profile
Doppler Wind Experiment	Measure winds, speed and direction
Nephelometer	Cloud structure Solid/liquid particles
Net-flux radiometer	Thermal/solar energy

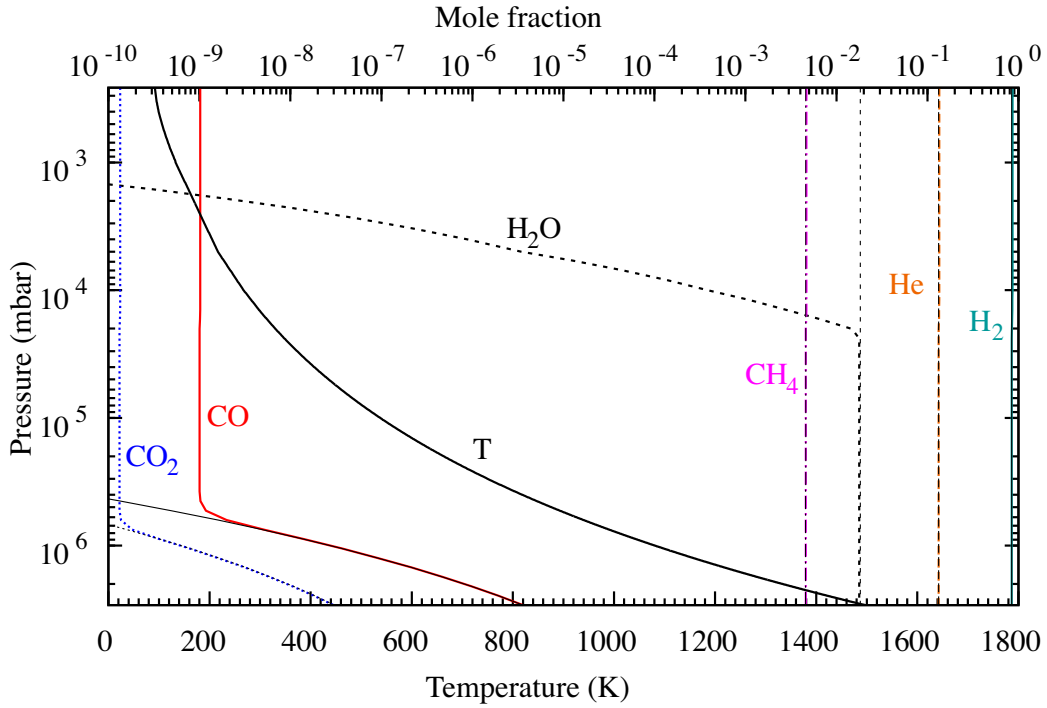


Figure 1: Mole fraction profiles in the troposphere of Saturn obtained with [Venot et al. \(2012\)](#)'s model, targeting the  $10^{-9}$  upper limit on the upper tropospheric CO mole fraction obtained by [Cavalié et al. \(2009\)](#). The temperature profile in the troposphere is shown in black solid line. Thermochemical equilibrium profiles are shown as black solid lines with the same layout as their corresponding species. The model parameters are: O/H= 21 times solar, C/H= 9 times solar, and  $K_{zz} = 10^9 \text{ cm}^2 \cdot \text{s}^{-1}$ . Condensation of  $\text{H}_2\text{O}$  occurs around the 20 bar level in this model.

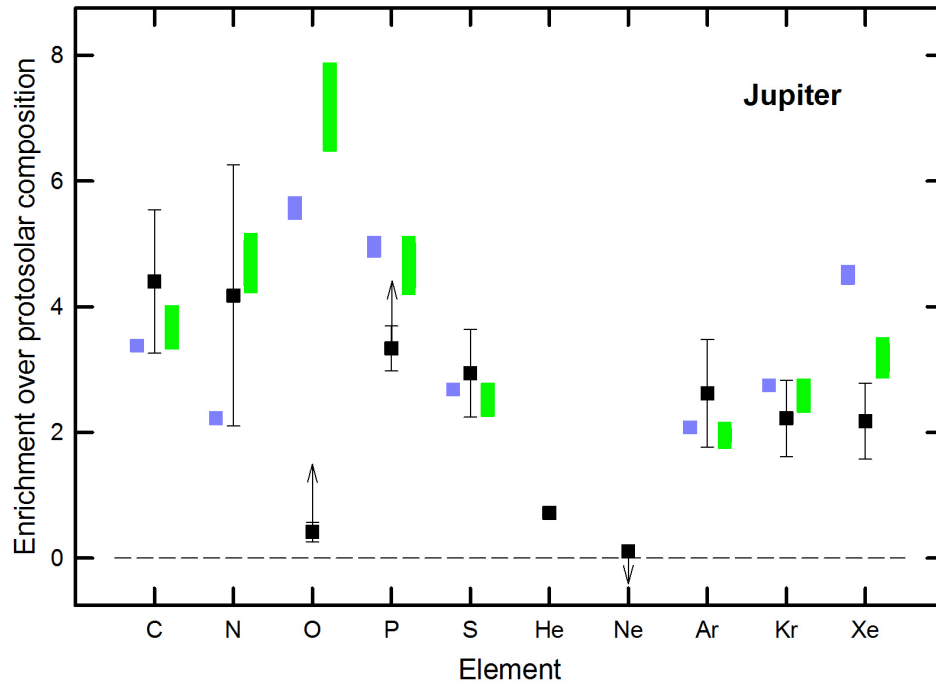


Figure 2: Ratio of Jovian to protosolar abundances. Black squares and black bars correspond to measurements and their associated uncertainties. Blue and green bars correspond to calculations assuming oxidizing and reducing conditions in the protosolar nebula, respectively (see text). Arrows pointing up correspond to the possibility that the measured oxygen and phosphorus abundances are lower than their bulk abundances, and arrow pointing down to the fact that the measured Ne abundance is an upper limit.

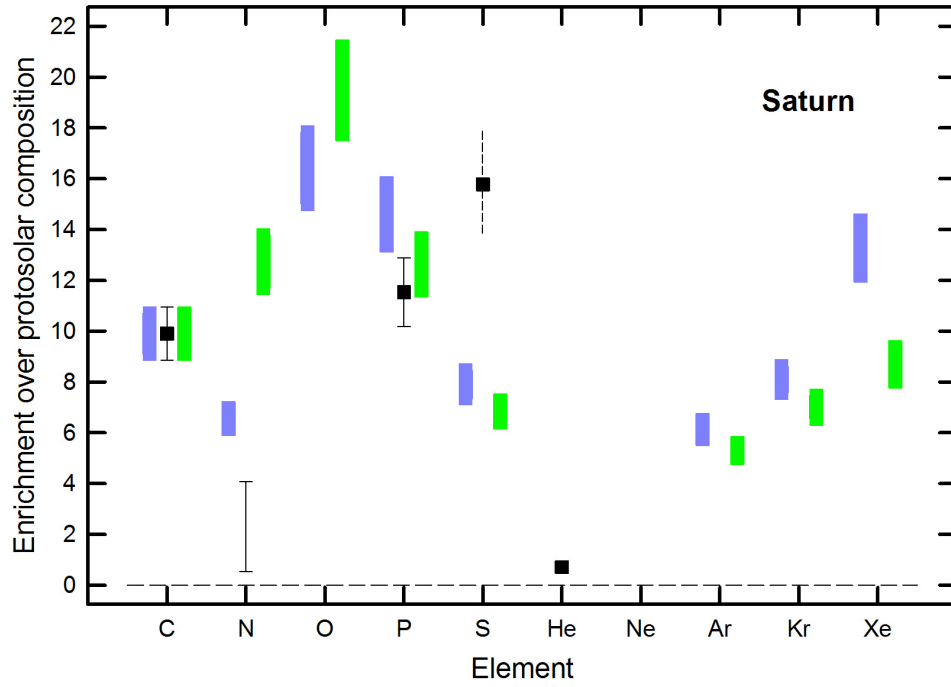


Figure 3: Ratio of Saturnian to protosolar abundances. Black squares and black bars correspond to measurements and their associated uncertainties. The O value measured in the troposphere would be close to zero on the utilized scale. Blue and green bars correspond to calculations assuming oxidizing and reducing conditions in the protosolar nebula, respectively (see text).